

RAPID COMMUNICATION

Dinosaur footprints in the Early Jurassic of Patagonia (Marifil Volcanic Complex, Argentina): biochronological and palaeobiogeographical inferences

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Abstract

A new dinosaurian track-bearing site, with tridactyl footprints from the Lower Jurassic (pre-middle Pliensbachian) volcanogenic and epiclastic rocks of the Marifil Volcanic Complex, Patagonia, Argentina, is presented and described. The best-preserved footprint, classified as cf. *Anomoepus*, confirms the utility of the *Anomoepus*-like tracks for the Early Jurassic biochronology. Palaeobiogeographically, this record supports the idea that the South American Early Jurassic dinosaur fauna presents elements of Pangaean distribution, and others with Gondwanan relationships with prevalent southern African affinities. Dinosaur records from South America between the Rhaetian and the Pliensbachian are very scarce, and this find contributes to the knowledge of early radiation and evolution of Dinosauria.

Keywords: *Anomoepus*-like tracks, pre-middle Pliensbachian, ichnology, South America, volcanogenic rock

1. Introduction

Dinosaurs originated, radiated and became the dominant vertebrate group in continental tetrapod communities of the world during most of the Mesozoic (Benton, 1983; Brusatte *et al.* 2008, 2010). An important part of their history is preserved in the sedimentary successions cropping out in Patagonia, Argentina (e.g. Casamiquela, 1964; Bonaparte & Vince, 1979; Salgado & Bonaparte, 1991; Coria & Salgado, 1995; Novas, 2009), ranging from the early stages after their origin in the Late Triassic up to their extinction in the uppermost Cretaceous rocks (see Novas, 2009, and references therein). Nevertheless, the history is far from being completely known and any new data will be crucial to understand both the early radiation of Dinosauria in Patagonia and the first steps of the Mesozoic tetrapod fauna that evolved in Gondwana.

During the 1950s, a new dinosaurian track-bearing layer was found in a flagstone quarry in the Lower Jurassic age (pre-middle Pliensbachian; see below) Marifil Volcanic Complex (MVC), in the SE of Río Negro province, Patagonia, Argentina. To date, there have been no discoveries of

vertebrate remains in the MVC. For the large temporal interval of *c.* 26 Ma from the Rhaetian (uppermost Triassic) to the Pliensbachian (middle Early Jurassic), dinosaur remains have been very scarce in South America (Barrett *et al.* 2008, 2014; Martínez, 2009; Apaldetti *et al.* 2011; Pol, Garrido & Cerda, 2011; Langer *et al.* 2014). This contrasts with the abundant information available from younger and older levels (e.g. Bonaparte, 1971; Báez & Marsicano, 2001; Arcucci, Marsicano & Caselli, 2004; Salgado & Gasparini, 2004; Langer, 2005; Rauhut *et al.* 2005; Pol & Powell, 2007; de Valais, 2011; Pol, Rauhut & Becerra, 2011; Pol & Rauhut, 2012).

This study aims to present, describe and analyse the new ichnological material found in the MVC. In addition, the biochronological and palaeobiogeographical significance of these dinosaur footprints is discussed in the context of the geological evolution and global tetrapod faunal composition of Patagonia and also Gondwana.

2. Geological setting

The MVC represents a large magmatic Mesozoic event in the eastern North Patagonian Massif, Argentina (Pankhurst *et al.* 1998, and references therein). It comprises a large volume of acidic (rhyodacites to rhyolites) ignimbrites with minor rhyolitic and andesitic lava flows, and sedimentary lenses interbedded within the acidic volcanic succession (Cortés, 1981). A variety of igneous rocks from the MVC have been dated by several radiometric methods (Rb–Sr, K–Ar, Ar–Ar and U–Pb), ranging from 221 to 165 Ma (Cortés, 1981; Pankhurst *et al.* 1998, 2000; Féraud *et al.* 1999, and references therein). Because of its overall rhyolitic composition and the proposed Jurassic age, the MVC has been included in the Chon Aike Large Silicic Igneous Province (Pankhurst *et al.* 2000).

In the study area, rhyolitic ignimbrites of 188 Ma (Rb–Sr age, in Pankhurst & Rapela, 1995) overlay a series of acidic to mesosilicic igneous and pyroclastic rocks and thin epiclastic lenses. The ichnosite is located 50 km SW of Sierra Grande, Río Negro province, in a farm owned by the Perdomo family (Fig. 1). The 30 m succession containing the flagstone quarry, which had fielded the track-bearing slabs, is dominated by pyroclastic acidic flows with a thin volcanogenic epiclastic lens (Fig. 2). The track-bearing slabs are composed of coarse-grained light-pinkish sandstone with

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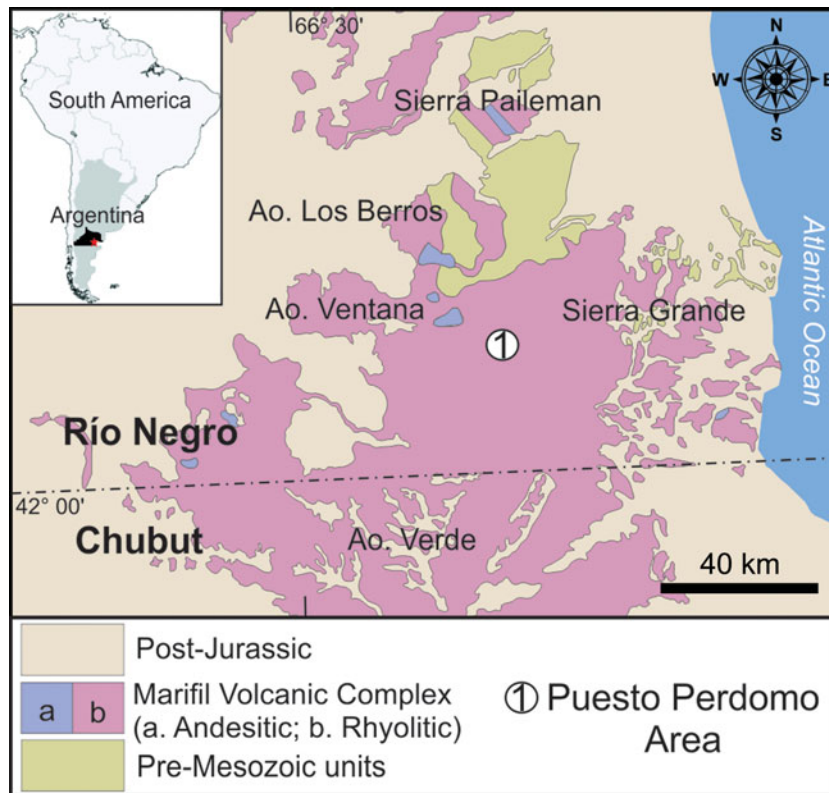


Figure 1. (Colour online) Local geological map with the location of Sierra Grande – Arroyo de la Ventana area and the new dinosaurian track-bearing site.

quartz, k-feldspar and pyroclastic material (ash and pumice fragments), which come from the base of the 1 m thick epiclastic lens in the analysed succession (Fig. 2). The coarse sandstones are interbedded with fine-grained brownish sandstone on which the dinosaur probably stepped, given that the tracks studied here are the infill of the original footprints. At the top of epiclastic lenses, plant remains attributable to equisetals (M. G. Passalia & A. Iglesias, pers. comm., 2015) were found. Flat lamination and ripple marks are also present at the epiclastic levels.

The described epiclastic rocks and dinosaur tracks are time-constrained by the andesitic rocks at the base of the MVC (221 Ma, Carnian, Late Triassic) and the upper acidic rhyolitic ignimbrites at the top (188 Ma, early Pliensbachian, Early Jurassic). Because of their acidic pyroclastic component precedent from the overlying rhyolitic complex, the epiclastic lenses are assigned to the Early Jurassic gap of the Marifil Volcanism (pre-middle Pliensbachian) (Cortés, 1981; Pankhurst & Rapela, 1995).

The thickness of the epiclastic succession and the observed palaeontological and sedimentological features suggest that the sedimentation occurred in a small fluvial system, where the sandstones correspond with the erosion of the ignimbrites infilling small palaeochannels by medium-energy currents. The volcanism of the MVC may have controlled the sedimentation and the development of the fluvial system.

3. Material and methods

The tracks were collected *in situ* by the Perdomo family in the 1950s, from a flagstone quarry; in the 2000s, the material was donated to the Museo Regional Provincial de Valcheta, Valcheta town, Río Negro province, and housed under the acronym MRPV. The specimens are six trace

fossils preserved as positive relief, in four sandstone slabs (Fig. 3). Their collection numbers are MRPV 427/P/13, 428/P/13, 429/P/13, 430/P/13.1, 430/P/13.2 and 430/P/13.3 (the last three specimens are in the same slab).

The ichnotaxonomic approaches to tridactyl footprints of Gierliński (1991), Olsen & Rainforth (2003) and Li *et al.* (2012) have been followed. Measurements and nomenclature are mainly based on the criteria of Leonardi (1987) and Haubold (1971). The measurements (Table 1) were: footprint length (FL), footprint width (FW), digit impression length (II, III and IV) and digit impression divarication angles (II–III, III–IV, II–IV). A further parameter is the ratio between the maximum height and the perpendicular transverse base of the anterior triangle (AT) formed by digit II, III and IV tip imprints (*sensu* Lockley, 2009, and references therein).

Photogrammetric models (Mallison & Wings, 2014) (Fig. 3) were obtained using Agisoft PhotoScan™ (version 0.8.5.1423) software (Grupo Aragosaurus, Universidad de Zaragoza License), and imported into Meshlab (version v1.3.3) and Paraview (version 3.14.1) software packages in which depth and contour line analysis was produced.

4. Results

The four pedal impressions are tridactyl, subsymmetric and mesaxonic. They are longer than wide and present the ‘heel’ impression almost directly or directly aligned with the axis of digit III impression.

MRPV 427/P/13 (Fig. 3a) is a natural cast, 131.3 mm long and 100.3 mm wide (length/width ratio: 1.31). The digit impressions are slender, longer than wide and are 55.5, 81 and 61 mm long for digits II, III and IV imprints, respectively. Claw and pad impressions are not evident. The divarication angles are II–III 31°, III–IV 51° and II–IV 68°. The AT is 0.51.

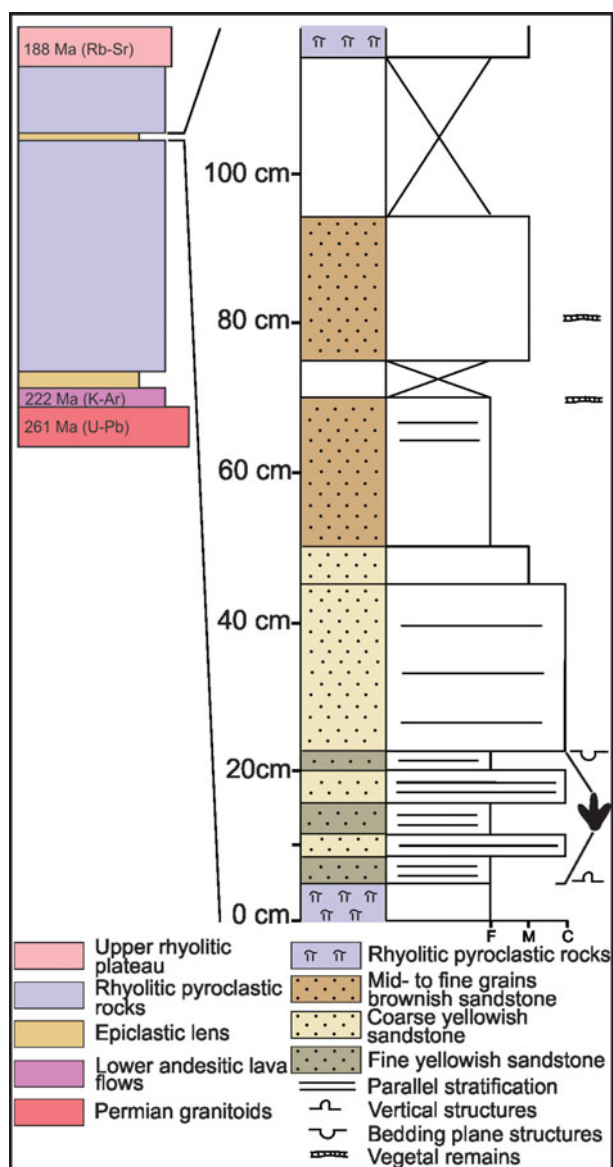


Figure 2. (Colour online) General, simplified, stratigraphic section of MVC and detailed stratigraphic section of Perdomo's quarry. Ages are based on references in the text.

MRPV 428/P/13 is a very irregular and poorly preserved natural cast, with a centimetric layer (undertrack cast) covering the footprint (Fig. 3b). The digit impressions are broad, wider than long. The track is about 172 cm long and 135 cm wide but we consider these measurements unrepresentative due to poorly-preservation.

MRPV 429/P/13 (Fig. 3c) is a natural cast that is 175 mm long and 127 mm wide (length/width ratio: 1.37). The digit impressions are slender and longer than wide. The digit III impression is 111 mm long. The other digit impressions are too poorly preserved to provide reliable measurements. The divarication angle II–IV is 52° . The AT is 0.53.

The three impressions of MRPV430/P/13 are preserved as natural casts. MRPV 430/P/13.1 (Fig. 3d–f) is the best preserved. It is a footprint 187 mm long and 128.3 mm wide (length/width ratio: 1.46). The digit impressions are slender, longer than wide and show clear digital pad imprints. The digit II impression is shifted anteriorly with respect to digits III and IV, displaying a characteristic posteromedial notch, indicating this is a right *pes*. The relative digit lengths are $III > IV > II$ (106, 81 and 71 mm). The metatarsophalangeal

pad trace is very clear, as well as the claw traces of digit II and III impressions, laterally and anteriorly directed respectively. The impression of digit IV projects slightly further than the digit II impression. The divarication angle II–IV is 58° (II–III is 19° , III–IV is 27°). The AT is 0.50. On the same slab, there are two ovoid traces positioned close to the tridactyl track MRPV 340/P/13.1 (Fig. 3d–f). One (430/P/13.2), of about 80 mm diameter, is located 100 mm on the left of the 'heel' impression, and the other (430/P/13.3), of about 50 mm diameter, is 70 mm in front of the anterior part of the digit III impression.

5. Discussion and conclusions

5.a. Ichnotaxonomy

The studied footprints display different kinds of preservation. With this in mind, the general and recurrent tridactyl shape shows that the tracks are all very alike and variation in the morphology may represent the product of taphonomical variability. Because of the observed preservational variants due to taphonomy, tracks were classified in different ichnotaxonomical levels, from the best-preserved track MPV 430/P/13.1, assigned to a higher ichnotaxonomical status, to the other three tracks. In fact, the poorly preserved MVP 427/P/13, 428/P/13 and 429/P/13 do not display sufficiently clear morphological details to undertake an ichnotaxonomical assignment with confidence.

The principal features shared by the footprints (i.e. tridactyl, roughly symmetrical, mesaxonic, longer than wide, 'heel' impression in line with the axis of digit III impression) are common in some theropod and ornithischian ichnotaxa, such as *Anomoepus* Hitchcock, 1848, *Ornithomimus* Sternberg, 1926, *Saurexalopus* Harris, 1997 or *Dinehichnus* Lockley *et al.* 1998, among others (Wright, 2004). MVP 428/P/13 presents broader digit impressions than the other three tracks, likely due to the natural-cast flattening phenomenon (Lockley & Xing, 2015). In the case of MVP 427/P/13, 428/P/13 and 429/P/13, their general features are approximately their whole description and they have no other peculiarities to relate them to a particular ichnotaxon. Therefore, we classify these footprints as indeterminate dinosaur footprints.

MPV 430/P/13.1 is the best-preserved specimen from the Perdomo site. It mainly differs from the typical Late Triassic–Early Jurassic theropod ichnotaxa *Eubrontes* Hitchcock, 1845, *Anchisauripus* Lull, 1904 and *Grallator* Hitchcock, 1858 because the former is roughly symmetrical and presents a metatarsophalangeal pad impression in line with the axis of the digit III impression, while the latter are asymmetrical with the metatarsophalangeal pad impression laterally located.

MPV 430/P/13.1 displays similar ichnotaxobases to some ichnogenera that belong to the ornithischian ichnofamily Anomoepodidae Lull, 1904 (i.e. symmetry, position and the shape of the metatarsophalangeal pad impression and digital pad impressions; Fig. 3d–f), such as *Anomoepus*, *Moyenisauropus* Ellenberger, 1970, and *Shenmuichnus* Li *et al.* 2012 (*Anomoepus*-like ichnotaxa in this work). These ichnogenera share several features (bipedal or quadrupedal trackways, pentadactyl hand prints and tridactyl or tetradactyl footprints), causing their ichnotaxonomy to be disputed. For instance, some authors have considered *Moyenisauropus* as a junior synonym of *Anomoepus* (Olsen & Galton, 1984; Olsen & Rainforth, 2003), while others argue that they are different ichnogenera (Gierliński, 1999; Lockley & Gierliński, 2006; Dalman & Weems, 2013). Li *et al.* (2012) suggest that *Shenmuichnus* has a lower heteropody

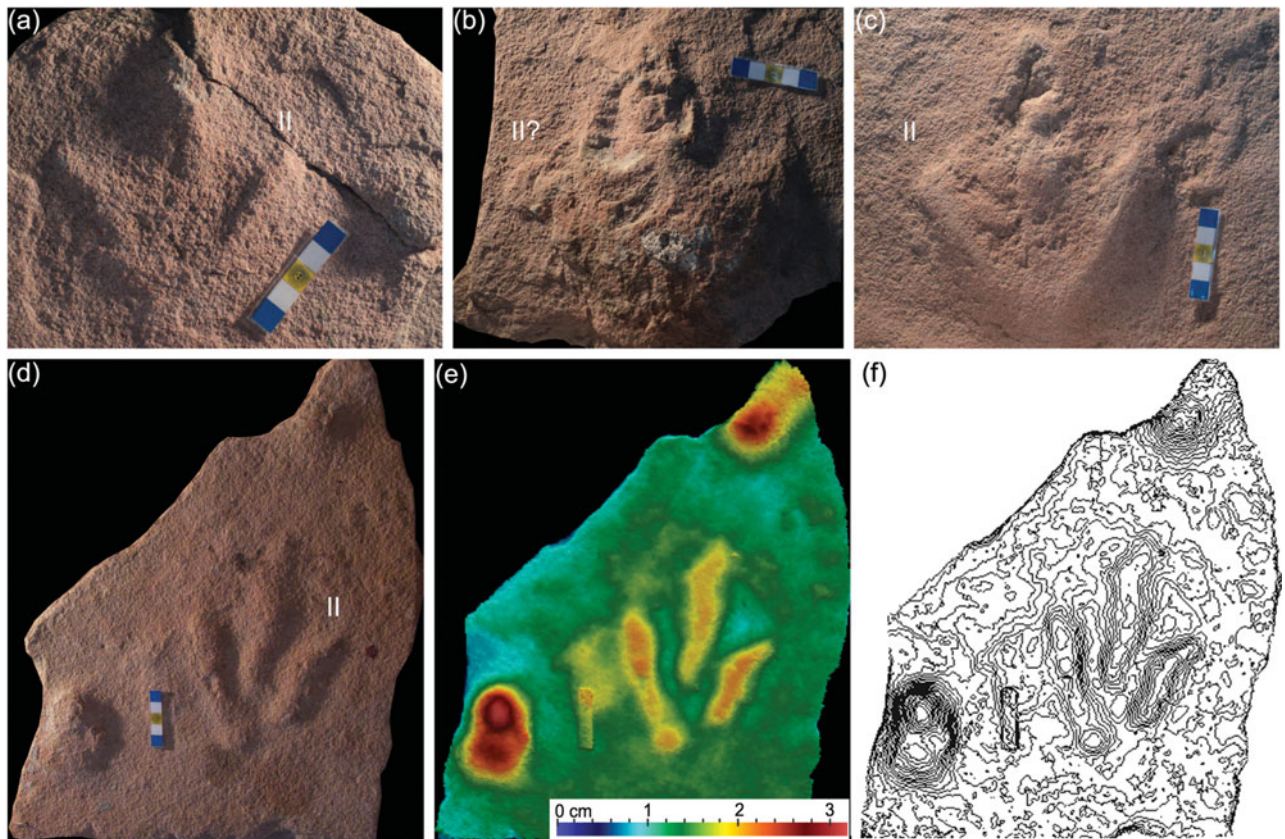


Figure 3. (Colour online) MVC dinosaur footprints. (a) MRPV 427/P/13; (b) MRPV 428/P/13; (c) MRPV 429/P/13; (d) MRPV 430/P/13; (e) false-coloured 3D depth analysis model of MRPV 430/P/13; and (f) contour line map with 1 mm of equidistance of MRPV 430/P/13. Scale bar: 5 cm. II: digit II impression.

from *Moyenisauropus* and *Anomoepus* and the lack of claw impressions. Gierliński (1991) described *Moyenisauropus* more robust than *Anomoepus* (as in *Shenmuichnus*; *sensu* Xing *et al.* 2016b) and with subequally lengthened digit impressions on the *manus* tracks. The MPV 430/P/13.1 footprint is gracile and preserves claw impressions, so we propose a relationship with *Anomoepus*. As noted above, the *manus* impression is important in the ichnotaxonomy of *Anomoepus*-like ichnotaxa. Close to MPV 430/P/13.1, there are two traces (MPV 430/P/13.2–3; Fig. 3d–f) that might be considered as *manus* impressions. Nevertheless, these impressions could also be interpreted as part of an undetermined, partial and poorly preserved distinctive trackway. Lockley & Gierliński (2006) suggested that *Anomoepus* is hard to identify with confidence unless both *manus* and *pes* impressions were found. Therefore, we classified MPV430/P/13.1 as cf. *Anomoepus* due to its similarity to this ichnotaxon but lack of a clear *manus* impression.

5.b. Trackmaker affinity and South American coetaneous dinosaur diversity

The *Anomoepus*-like tracks have been related to ornithischian trackmakers by several authors (e.g. Lull, 1904; Haubold, 1971; Olsen & Galton, 1984; Gierliński, 1991; Olsen & Rainforth, 2003). A criterion has been the presence of the dinosaurian *pes* and the pentadactyl *manus* impressions (Olsen & Rainforth, 2003). These authors suggest that the *manus* track lacks enlarged digit I, II and III impressions, which is related to the manual phalangeal formula of Ornithischia. Within this clade, basal members of ornithischians, ornithopods or thyreophorans have been cited as possible

trackmakers (e.g. Thulborn, 1990; Gierliński, 1999; Olsen & Rainforth, 2003; Li *et al.* 2012). Therefore, Olsen & Rainforth (2003) suggested that the producer of *Anomoepus* was a relatively small, gracile, facultatively bipedal ornithischian.

As stated above, dinosaur remains are very scarce in the Rhaetian–Pliensbachian of South America (Barrett *et al.* 2008, 2014; Martínez, 2009; Apaldetti *et al.* 2011; Pol, Garrido & Cerda, 2011; Langer *et al.* 2014) and the studied tracks represent the first ichnological record from this time interval in the region. Moreover, if we are correct in our appraisal that the trackmaker of MPV 430/P/13.1 was an ornithischian, then the studied tracks would also represent the second evidence of the occurrence of this clade in South America during the same time interval (the first is *Laquintasaura venezuelae* Barrett *et al.* 2014, lowermost Hettangian La Quinta Formation, Venezuela). This find could confirm the presence of ornithischians in Patagonia between the Norian cf. *Heterodontosaurus* sp. from the Laguna Colorado Formation (Báez & Marsicano, 2001), and the Toarcian *Manidens condoriensis* Pol, Rauhut & Becerra, 2011 and Heterodontosauridae indet. (Becerra *et al.* 2016) from the Cañadón Asfalto Formation.

Prior to this study, only one Patagonian dinosaur had ever been documented from the Rhaetian–Pliensbachian lapse: the basal sauropodomorph *Leoneerasaurus* Pol, Garrido & Cerda, 2011, from the Sinemurian–Pliensbachian Las Leoneras Formation (Cañadón Asfalto basin, Chubut province; age *sensu* Cúneo *et al.* 2013). The scarcity of vertebrate fossils from the Rhaetian–Pliensbachian time interval strongly contrasts with the abundant known fossil-bearing horizons from older (e.g. Casamiquela, 1964; Báez & Marsicano, 2001; Pol & Powell, 2007) and younger (e.g.

Table 1. Measurements of MVC dinosaur footprints

	FL (mm)	FW (mm)	II (mm)	III (mm)	IV (mm)	II–III (°)	III–IV (°)	II–IV (°)	AT
MRPV 427/P/13	131	100	55.5	81	61	31	51	68	0.51
MRPV 428/P/13	172?	135?							
MRPV 429/P/13	175	127		111				52	0.53
MRPV 430/P/13.1	187	128	71	106	81	19	27	58	0.50

Abbreviations: FL, footprint length; FW, footprint width; II: digit II length; III: digit III length; IV: digit IV length; II–III: angle between digits II and III; III–IV: angle between digits III and IV; II–IV: angle between digits II and IV; AT: anterior triangle. AT is an index and therefore dimensionless.

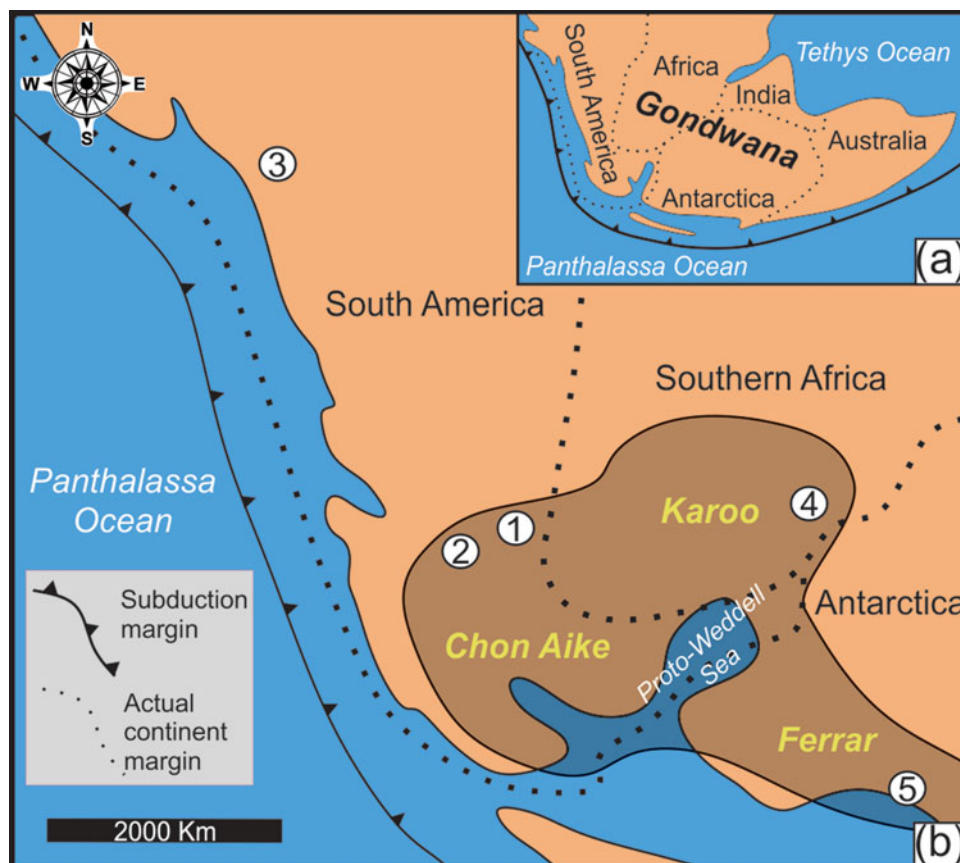


Figure 4. (Colour online) Palaeogeographic reconstruction of the SW Gondwana/Pangaea in the Late Triassic / Early Jurassic (based on Pankhurst *et al.* 2000; Golonka, 2007). (a) Regional view of the Gondwana landmass. (b) Detail of the reconstruction showing the Early Jurassic large igneous provinces of South Gondwana (Chon Aike, Karoo and Ferrar) and the distribution of the fossil localities mentioned in the text. 1. MVC tracksite; 2. Las Leoneras Formation; 3. Cañón del Colorado and Balde de Leyes formations; 4. Elliot Formation; and 5. Hanson Formation.

Salgado & Gasparini, 2004; Rauhut & López-Arbarello, 2008; de Valais, 2011; Pol, Rauhut & Becerra, 2011) deposits.

From out of Patagonia, there are only four records for the Rhaetian–Pliensbachian time interval in South America: (1) from Argentina, two basal sauropodomorphs, *Adeopapposaurus magnai* Martínez, 2009, and *Leyesaurus marayensis* Apaldetti *et al.* 2011, were defined from the Lower Jurassic Cañón del Colorado Formation (Martínez, 2009) and Balde de Leyes Formation (Apaldetti *et al.* 2011; Colombi *et al.* 2015), respectively; and (2) from Venezuela, the basal ornithischian *Laquintasaura venezuelae* and the theropod *Tachiraptor admirabilis* Langer *et al.* 2014, have been defined from the lowermost Hettangian La Quinta Formation (Barrett *et al.* 2008, 2014; Langer *et al.* 2014).

In view of the above, the dinosaur footprints studied herein are an important finding because they provide new

and valuable information about the scarce dinosaur record from South America during the uppermost Late Triassic–Early Jurassic. Furthermore, these tracks are the first vertebrate ichnological remains in this continent for the Rhaetian–Pliensbachian interval.

5.c. Biochronological and palaeobiogeographical inferences

Tetrapod footprints provide important data on the vertebrate record, both in space and in time distribution (Lucas, 2007). The footprints related to the *Anomoepus*-like ichnotaxa present a widespread geographical distribution and a particular temporal occurrence. *Anomoepus*-like footprints have been identified in the Early Jurassic from: (1) North America: USA (e.g. Olsen & Rainforth, 2003; Lockley &

Gierliński, 2006; Dalman & Weems, 2013); (2) Europe: Poland (Gierliński, 1991) and Italy (Avanzini, Gierliński & Leonardi, 2001); (3) Asia: China (e.g. Lockley & Matsukawa, 2009; Li *et al.* 2012; Xing *et al.* 2016b); (4) Oceania: Australia (e.g. Thulborn, 1994); and (5) Africa: Lesotho (Ellenberguer, 1970).

Other tracks related to the *Anomoepus*-like ichnotaxa from the Late Triassic of Poland (Niedźwiedzki, 2011) and the USA (Baird, 1964), and from the Middle Jurassic of China (Xing *et al.* 2015, 2016a) and Morocco (Belvedere *et al.* 2011) have been published. However, their ichnotaxonomical affinity or the proposed ages of the tracksites were questioned. Lockley & Gierliński (2006) suggested that the Late Triassic tracks from the USA classified as ?*Anomoepus* isp. by Baird (1964) are indeed chirotheriid tracks. The Poland tracks cf. *Anomoepus* isp. are partly eroded and slightly deformed (Niedźwiedzki, 2011), so it is difficult to relate them with confidence to this ichnogenus. The Middle Jurassic *Anomoepus*-like tracks of Morocco (Belvedere *et al.* 2011) and the Henan province of China (Xing *et al.* 2016a) are similar in shape. These tracks have very slender digit impressions and narrow metatarsophalangeal pad impressions, being closer to an avian-like ichnotaxon than to *Anomoepus*. The Middle Jurassic tracks of *Shensipus tungchuanensis* Young, 1966, were first related to theropods (Young, 1966; Lockley *et al.* 2013). Recently, Xing *et al.* (2015) proposed the new combination *Anomoepus tungchuanensis*. Actually, the specimens are lost (*sensu* Xing *et al.* 2015), and according to the original photographs they are poorly preserved (e.g. thin layer infill tracks, very shallow anterior surface; see Young, 1966), so a confident conclusion is not possible. The *Anomoepus* tracks from Shaanxi province, also in China (Xing *et al.* 2015), originally determined as Middle Jurassic in age, have recently been included in Lower Jurassic layers based on detailed stratigraphic work (Wang *et al.* 2016).

The age of MVC footprints is consistent with the known temporal distribution of *Anomoepus*-like tracks and may represent a spatially near-global biostratigraphic occurrence (Early Jurassic biochron) of this ichnotaxon (see Haubold, 1986; Lucas, 2007). In addition, this material represents the unique *Anomoepus*-like tracks from South America, increasing its record almost worldwide (except in Antarctic rocks).

The palaeobiogeographic connections between the vertebrate Gondwanan palaeofaunas during the Late Triassic–Early Jurassic have been widely recognized (e.g. Yates, 2003; Langer, 2005; Pol & Powell, 2007; Bittencourt & Langer, 2011). Similarities have previously been noted among the basal sauropodomorphs from the Norian Los Colorados Formation, Argentina (age *sensu* Kent *et al.* 2014), and the Norian Caturrita Formation, Brazil (age *sensu* Langer & Ferigolo, 2013), and the palaeofauna from the Norian–Rhaetian lower Elliot Formation, southern Africa (age *sensu* Knoll, 2005). These similarities support the hypothesis of a palaeofaunal interchange between the southern African and South American tetrapods during the Late Triassic. The Early Jurassic fauna of South America also presents phylogenetic affinities with the upper Elliot Formation, southern Africa (Rauhut & López-Arbarello, 2008; Martínez, 2009; Apaldetti *et al.* 2011; Pol, Garrido & Cerda, 2011; Sereno, 2012), although relationships with other Gondwanan (Antarctica) and Laurasian (China) zones have also been identified (see Smith & Pol, 2007; Rauhut & López-Arbarello, 2008; Apaldetti *et al.* 2011). This happens with the Early Jurassic footprints studied herein which are close to *Moyenisauripus* ichnotaxon from the southern African upper Elliot Formation (*sensu* Ellenberger, 1970). Nevertheless, MPV 430/P/13.1 is also similar to some Early Jurassic *Anomoepus*-like tracks found in other places of

Gondwana (Australia; Thulborn, 1994) and Laurasia (North America, Europe and Asia; e.g. Gierliński, 1991; Avanzini, Gierliński & Leonardi, 2001; Olsen & Rainforth, 2003; Lockley & Gierliński, 2006; Lockley & Matsukawa, 2009; Li *et al.* 2012). This idea is consistent with the Jurassic palaeoflora from Patagonia that is comparable with the Antarctica record, but present Pangaeian relationships as well (Wilf *et al.* 2013).

The break-up of Pangaea initiated during the Early Triassic (see Golonka, 2007, and references therein). Nevertheless, South America has remained connected to almost all the landmass of Pangaea through the Jurassic (Wilf *et al.* 2013). As well, Rapela *et al.* (2005) suggested Permian–Triassic proximity between Patagonia and southern Africa according to their palaeoflora record (Archangelsky, 1990; Artabe, Morel & Spalletti, 2003). Additionally, the uppermost Early Jurassic magmatism of Chon Aike (in which is included the MVC), Ferrar and Karoo large igneous provinces has been correlated by many authors and associated with a mantle plume precursor of the Weddell Sea opening and the separation of Gondwanan terranes (Elliot & Fleming, 2000; Pankhurst *et al.* 2000; Rapela *et al.* 2005). This supports the idea that the current Patagonia, Africa, Antarctica and others areas of Pangaea were connected during the Early Jurassic, at least until the Weddell Sea development. Therefore, it is understandable that the Early Jurassic palaeofauna from southern South America and especially Patagonia presents a heterogeneous composition, with elements of Pangaeian distribution, and others with Gondwanan relationships with prevalent southern African affinities.

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