SPACE-WEIGHTED SEISMIC ATTENUATION MULTI-FREQUENCY TOMOGRAPHY AT DECEPTION ISLAND VOLCANO (ANTARTICA)

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Theory and Method

- Analysis
- Results

DiscussionsOutlooks

THEORY 1/2

<u>The "Japanese School"</u> energy variations in the heterogeneous Earth



Haruo Sato · Michael C. Fehler Takuto Maeda

Seismic Wave Propagation and Scattering in the Heterogeneous Earth

Deringer

Second Edition

<u>The "framework":</u> building forward models with Radiative Transfer Theory



Seismic attenuation tomography in volcanoes

Mount St. Helens



Campi Flegrei



THEORY 2/2



- Applying the theory of radiative transfer, sensitivity kernels can be obtained by modeling scattering intensities at different lapse times;
- Use coda-wave "scattered" information;
- Absorption + scattering → Coda Quality Factor (Qc);
- Qc is measured from the decay of coda intensity versus lapse time.



Does not work at all frequencies and scales at the same way: we need results in different frequency bands.

METHOD 1/3

SENSITIVITY KERNELS FOR CODA IMAGING: REGIONAL/CONTINENTAL SCALES (MAYOR ET AL. 2016)



Linear relation between
the spatial variation of the inelastic quality factor Qi(x) and Qc:
$$Q_c^{-1}(\mathbf{R}, \mathbf{S}) \approx \frac{\int Q_i^{-1}(\mathbf{x}) K_a(\mathbf{R}, \mathbf{S}; \mathbf{x}, t) d\mathbf{x}}{t}$$

"Ka", the absorption sensitivity kernel, which depends on the position of the source (S) and receiver (R) and the lapse-time t in the coda.

METHOD 2/3

KERNELS-BASED IMAGING IN VOLCANOES

(DEL PEZZO ET AL. 2016)



Inhomogeneous scattering properties of volcanoes to map shape and dimensions of hot reservoirs and plumbing systems.

The **forward model** is built with Monte Carlo simulations of Radiative Transfer Theory equations.

Sensitivity is generally **maximum at source and receiver**, with wider illumination than in ray-dependent tomography.

METHOD 3/3

KERNELS-BASED IMAGING IN VOLCANOES (INVERSION MODEL) (DE SIENA ET AL. 2017)

The main assumption is that total coda attenuation is caused by the medium comprising the inversion grid;

The weighting functions provide the rows of the inversion matrix at the nodes after normalization for the total weight relative to the source-receiver pair;





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4520000

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The output is generated from the contribution of all nodes of the grid to the single-station Qc measurement

DATA ANALYSIS



DATA ANALYSIS



- 1. Zeros percentage;
- 2. Mean amplitude Noise windows;
- 3. Mean signal amplitude;
- 4. Signal to Noise ratio;
- 5. Mean amplitude Coda windows;
- 6. Mean amplitude max power;
- 7. Mean frequencies at max power;
- 8. Cross-correlation mean values;
- 9. Sum of the ampl. values > 500;
- 10. Amplitude "jumps";
- 11. Mean Correlation coefficients; 12. Deviation Standard STA/LTA;

DATA ANALYSIS

| Zeros% | Amp Noise w | Amp Signal | S/N | Amp Coda w | MaxPow Amp | Freq MaxPow | Xcorr | Ampl>500 | Jumps | CorrCoe | DevStStaLta | Num. Events |
|-------------|-------------|-------------|------------|-------------|-------------|-------------|-------------|-------------|------------|-------------|-------------|-------------|
| 24,14472711 | 41,6522254 | 169,4206531 | 658,844695 | 86,36893258 | 1501055,616 | 11,30141498 | 2441546,661 | 49,08761031 | 53,3617604 | 0,637436169 | 1,12861352 | 20283 |
| 8,899726155 | 51,15679938 | 198,735286 | 35,9273832 | 111,5881513 | 1919600,057 | 10,81084024 | 3043020,852 | 59,4691424 | 65,6010488 | 0,558934996 | 0,81293642 | 14972 |
| 8,839893171 | 45,3949275 | 195,5547233 | 37,0140766 | 105,7499993 | 1778461,021 | 11,00276994 | 2880608,398 | 53,41411675 | 62,3985658 | 0,608990486 | 0,85451556 | 13105 |
| 8,914667511 | 18,37416775 | 168,6745146 | 40,5085317 | 80,80920191 | 1197581,724 | 11,78 | 2140024,209 | 30,20443319 | 48,1508695 | 0,798962086 | 1,03212687 | 7895 |
| 8,595400222 | 18,07422623 | 169,4120495 | 35,7345383 | 81,03120393 | 1189359,401 | 11,83091996 | 2142433,739 | 29,16050584 | 48,3483346 | 0,800404036 | 1,02829843 | 7197 |

Itera**lierdienali@ian449a%Btrapidili@i**€ihts≤ 0.6" 20283 → 7197

Filtering in five frequency bands, centered at 6, 9, 12, 15 and 18 Hz; 4 seconds of Coda windows (starting from the main peak); Grid resolution of 1 and 2 km.

RESULTS (1km resolution)

lte**Caritgien cali (lads1**særle)



6 Hz

15 Hz

RESULTS (1km res.)

Original dataset







Iteration: 4









Tikhonov reg. (1km res. – 15 Hz)

0.118





11.8





RESULTS (2km resolution)

lte**Caritgien cali (lads1**særle)



6 Hz

15 Hz

RESULTS (2km res.)

Original dataset







Iteration: 4







Tikhonov reg. (2km res. – 15 Hz)

0.155





15.5









| Q_{c}^{-1} | | | |
|------------|-------|-------|------|
| 0 | 0.004 | 0.024 | 0.11 |

Aa# = High attenuation Ba# = Low attenuation

We consider the high attenuation areas as rocks caracterized by hot fluids or magmatic batches and the low attenuation ones as cooling magmatic body.





Ba1

Aa# = High attenuation Ba# = Low attenuation

We consider the high attenuation areas as rocks caracterized by hot fluids or magmatic batches and the low attenuation ones as cooling magmatic body.

DISCUSSIONS

Comparison with previous studies: • Geology • Tectonics • Geochemistry • Geophysics • Volcanic hazard

Comparison with previous geological studies











Hawkes et al. 1961 – Fig. 3 Volcanic centres of the «pre-caldera» group.





Ba2





Hawkes et al. 1961 – Fig. 6 Volcanic centres of the «Neptune bellow» group.

1 Km













Comparison with previous tectonics studies





Paredes et al. 2006 – Fig. 4 Spatial distribution of morpholineaments and tectonic zoning.











Lopes et al. 2015 – Fig. 4 Main structural alignments that control the morphology











Lopes et al. 2015 – Fig. 4 Main structural alignments that control the morphology

Comparison with previous geochemical studies



Somoza et al. 2004 – Fig. 8 As and Mn distribution













Comparison with previous geophysical studies











Berrocoso et al. 2008 – Fig. 10 Deformation processes during 91/00 (NNW-SSE) and 02/03 (NE-SW)









Catalan et al. 2014 – Fig. 4a Positive magnetic anomaly close by the Low Attenuation area (Ba1)

Ba2

1 Km



Zandomeneghi et al. 2009 – Fig. 6 Velocity tomography model (1km depth)

Qi







Prudencio et al. 2013 The high attenuation areas (yellow to red) matches our anomalies.







Del Pezzo et al. 2016 Middle-point weighting functions to the single-station measurements

Comparison with previous volcanic hazard studies



Bartolini et al. 2014 – Fig. 1A Simplified regional tectonic map.



Bartolini et al. 2014 – Fig. 1B



Bartolini et al. 2014 – Fig. 7 Susceptibility map of future eruptions calculated with QVAST



Bartolini et al. 2014 – Fig. 12 Qualitative hazard map





Ba2

1 Km





Berrocoso et al. 2006 – Fig. 5.10-6 Map of natural hazards





Ba1





Smellie et al. 2002 – Fig. 6.3 Suggested escape route. The extraction point H (SE) spatially matches the high attenuation area Aa3.

1 Km

Ba₂

OUTLOOKS (1) at Deception Island Volcano

- 1. Repeat the analysis using different Coda Windows;
- 2. 4D analysis;

OUTLOOKS (2) Method application

- 1. Apply to other volcanoes the aforementioned combination between Qc-Kernel and GIS analysis;
- 2. Apply the dataset cleaning procedure to other big and high-dimensional data.

