

# Maastrichtian-Danian Northpatagonian rocky shore, Argentina

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## ABSTRACT

The Atlantic Maastrichtian-Danian (K/P) transgression over northern-central extra-Andean Patagonia (Argentina) covered both Mesozoic sedimentary basins and the Northpatagonian Massif (NPM). The flooding of the NPM resulted in a regional unconformity/nonconformity (70,000 km<sup>2</sup>) between the pre-Cretaceous basement (Jurassic Chon Aike Silicic Large Igneous Province and igneous-metamorphic Paleozoic basement) and the K/P marine transgressive record, constituting the widest known ancient rocky shore of South America (Northpatagonian rocky shore).

The transgressive stratigraphic record over the basement is mainly composed of isolated carbonate bioclastic deposits up to 40 m thick with predominance of bivalves, echinoderms, bryozoan, coralline red algae, and foraminifera skeletal remains; we interpreted these bioclastic near-shore deposits as rocky shore associations. The hard substrate, irregular seacoast and low accommodation space over the NPM provided a preferential ecological niche for encrusting biota (i.e., Oyster reefs) during the K/P epicontinental flooding. The colonization of “engineer ecosystems” organisms over thousands of square kilometers probably enhanced the coastal biodiversity. The K/P Northpatagonian rocky shore favored the conformation of a short-lived, transgressive, cool-water carbonate factory in the south-eastern extreme of South America.

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## 1. Introduction

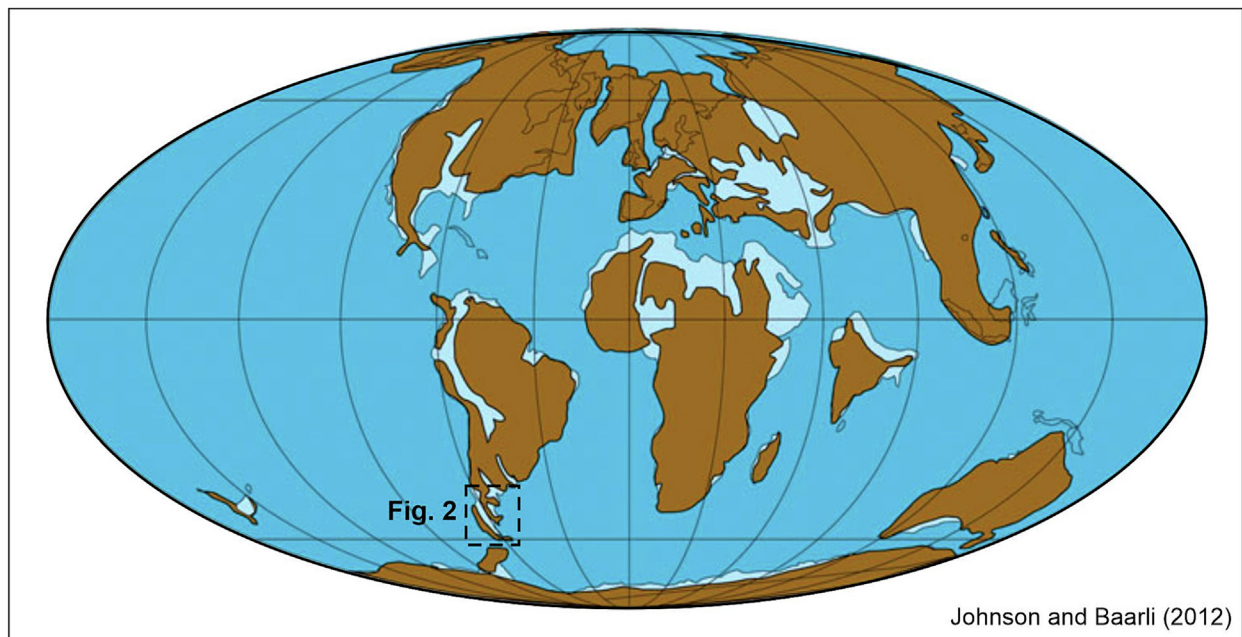
Coastal marine environments have been profusely studied because they represent a dynamic interplay between physic/biological and continental/marine processes. Although >70 % of the world's coastlines are rocky shores (Nyberg and Howell, 2016), ancient examples of this environment have been somewhat neglected by paleontologists and geologists until the 90s (Johnson, 1988a, 1988b, 1992). Geomorphologically, rocky shores often contain coastal cliffs, sea caves, bays, boulder shores and islands (i.e., Semeniuk and Johnson, 1985; Johnson et al., 1996; de Gibert et al., 1998; Betzler et al., 2000; Moura et al., 2006; Andriani and Walsh, 2007; among others). Geologically, ancient rocky shores are unconformities or nonconformities between hard-substrates and shallow marine deposits (i.e., Surlyk and Christensen, 1974; Johnson, 1988a, 1988b, 1992; Johnson and McKerrow, 1995; Johnson et al., 1996; Surlyk, 1997; Johnson and Baarli, 1999, 2012; Felton, 2002; Desrochers, 2006; Surlyk and Sørensen, 2010; Sørensen and Surlyk,

2010, 2011, 2013, 2015; Sørensen et al., 2011, 2012; Schröder et al., 2019; Puig López et al., 2023; among others). Biologically, rocky shores are very well-studied dynamic marine ecosystems where the substrate provides varied habitats for encrusting biota (i.e., Lewis, 1964, 1986; Dayton, 1975; Paine, 1994; Johnson and Baarli, 1999; Cebrián et al., 2000; Reise, 2001; Spencer and Viles, 2002; Johnson, 2006; Buatois and Encinas, 2011; Chappuis et al., 2014; Pineda-Salgado et al., 2015; McAfee et al., 2017, 2022, and cites therein). In this way, rocky shores and bioclastic deposits are a very frequent configuration during transgressive events because combine shallow marine conditions and hard-substrate (i.e., Webb, 1994; Jia-yu and Johnson, 1996; Sanders, 1998; de Gibert et al., 1998, 2012; Betzler et al., 2000; Larsen et al., 2003; D'Alessandro et al., 2004; Shepard, 2006; Felton et al., 2006; Buatois and Encinas, 2011; Bover-Arnal et al., 2011; Brlek et al., 2018; Sanders et al., 2019).

The worldwide cartography of marine shores for the last 250 Ma (Smith et al., 1994; Johnson, 2006; Johnson and Baarli, 2012) shows a maximum during Paleocene times with major seaways that deeply penetrated the North America, South America, Africa, and Europe (Fig. 1). In South America, the Upper Cretaceous sequences display progressive overstepping of older terranes during the epeiric flooding (i.e., Biddle

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**Fig. 1.** Global Early Paleocene paleogeographic configuration displaying major seaways penetrating most of the continents during a higher stand in sea level (taken from Johnson and Baarli, 2012, based on Smith et al., 1994). Location of Fig. 2 in the southern extreme of South America.

et al., 1986; Uliana and Biddle, 1988), producing a spectacular increase in the size of areas under marine influence of Argentina (i.e., Malumián et al., 1983; Uliana and Biddle, 1988).

This first Atlantic transgression widely flooded Andean and extra-Andean regions of Patagonia (southern Argentina) during the Maastrichtian-Danian (Fig. 2), representing the widest Cenozoic continental flooding in this region (Malumián and Caramés, 1995). Most of the regional contributions about the Maastrichtian-Danian transgression over central-northern Patagonia were dedicated to lithological/paleoenvironmental characterizations (i.e., Feruglio, 1949; Spalletti et al., 1993; Malumián, 1999), paleontologic/biostratigraphic studies (i.e., Burckhardt, 1901; Feruglio, 1936, 1949; Bertels, 1973, 1975, 1995; Camacho, 1992; Casadio, 1998; del Río, 2004, 2021; Guler et al., 2019; Brezina et al., 2021), paleobiogeographic/paleogeographic reconstructions (i.e., Malumián et al., 1983; Malumián and Nández, 2011; Aguirre-Urreta et al., 2011), diagenetic studies (Matheos and Tunik, 1998; Matheos et al., 2003) or subsidence analysis (i.e., Gianni et al., 2018a, 2018b; Dávila et al., 2019). Although the K/P transgressive record was described as lying over Jurassic volcanic rocks in the NPM (i.e., Feruglio, 1936, 1949; Simpson, 1941; Camacho, 1992; Aragón et al., 2014; Aguilera et al., 2014, among others) conforming a carbonate platform (Spalletti et al., 1993), just locally was interpreted as a rocky shore (Foix et al., 2015).

This work aims to characterize the K/P transgressive rocky shore and their associated deposits over the NPM, to estimate their areal extension, to approximate the main control factors in its development and contributing to understand the stratigraphic evolution of the extra-Andean Patagonia (Argentina).

## 2. Geological and paleontological framework

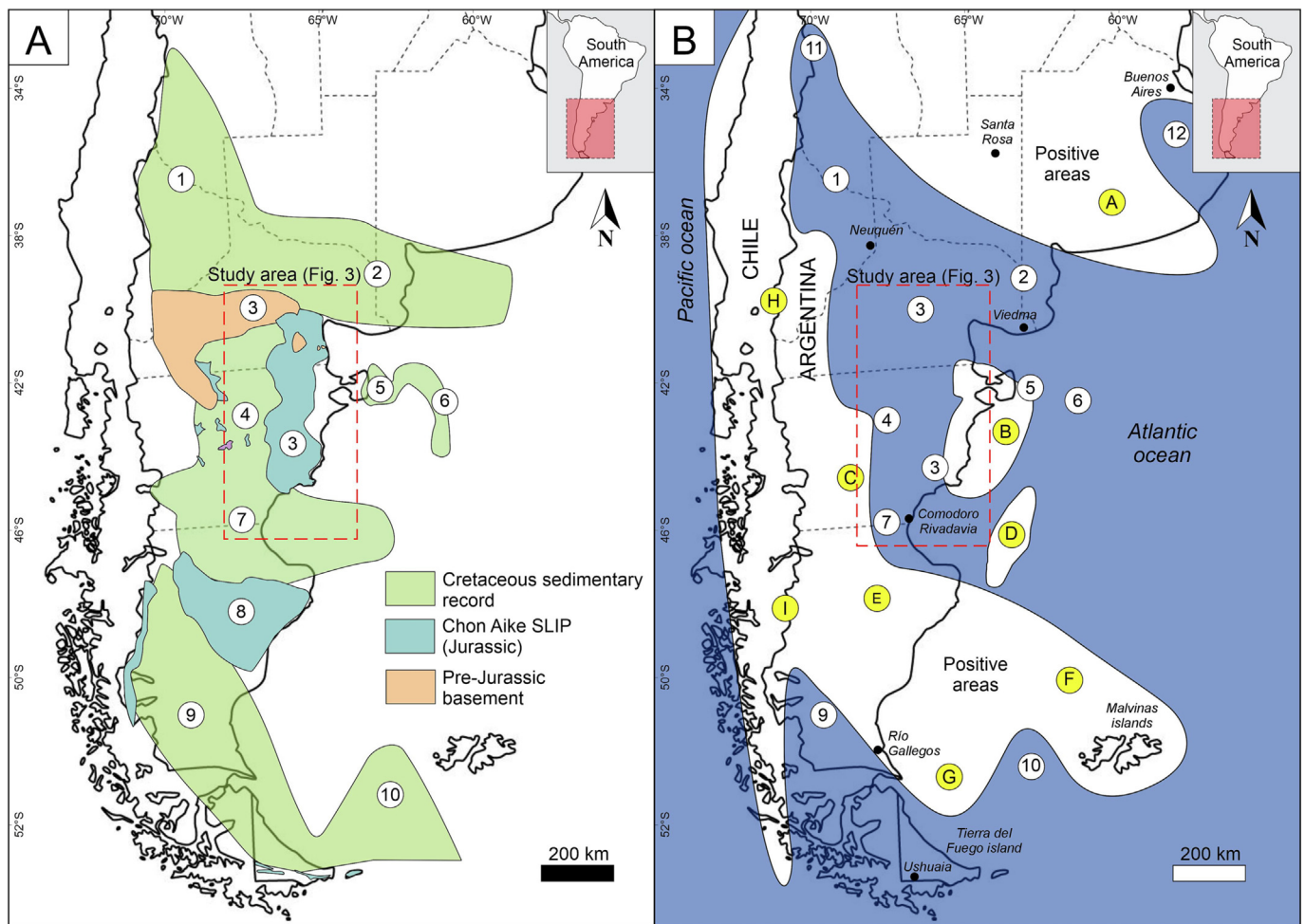
The break-up of Gondwana resulted in different Mesozoic volcanic provinces developed over pre-Mesozoic basements in South America, including the Chon Aike SLIP in Patagonia (Kay et al., 1989; Pankhurst et al., 1998; Bryan and Ernst, 2008; Bryan et al., 2010; Bryan and Ferrari, 2013; Lovecchio et al., 2020; Navarrete, 2021; Navarrete et al., 2021; among others). This SLIP covers 675.000 km<sup>2</sup> (Pankhurst et al., 1998), constituting one of the most important geologic features of the

region (Fig. 2A) with several lithostratigraphic units in different geological settings. This regional volcanic record also constitutes the synrift stage in several Jurassic-Cretaceous Patagonian sedimentary basins (i.e., Uliana et al., 1990; Franzese et al., 2003; Lovecchio et al., 2020) such as Golfo San Jorge (Clavijo, 1986; Fitzgerald et al., 1990; Figari et al., 1999; Sylwan, 2001; Figari and Hechem, 2021), Cañadón Asfalto (Figari and Courtade, 1993; Figari et al., 2015; Figari and Hechem, 2021), Austral-Magallanes (Franzese et al., 2003; Fildani et al., 2008; Sachse et al., 2016), Valdés (Continanzia et al., 2011) and San Julián (Homocv and Constantini, 2001).

Simplified, the basement of the NPM is mainly characterized by three groups of rocks: 1) Paleozoic igneous-metamorphic rocks such as Mina Gonzalito Complex, Yaminué Complex, Nahuel Niyeu Formation, El Jagüelito Formation and Punta Sierra Plutonic Complex (i.e., González et al., 2018, 2020), 2) Triassic volcanic/volcaniclastic suites represented by Los Menucos Complex (Cucchi et al., 2001; Labudía and Bjerg, 2001; Lema et al., 2008), and 3) extensive outcrops of the Jurassic Chon Aike SLIP represented by the Marifil Volcanic Complex (MVC), mainly composed of rhyolites, ignimbrites, and breccias (Pankhurst et al., 1998; Márquez et al., 2010; Strazzere et al., 2022). The NPM was strongly eroded from the Middle Jurassic to the Late Cretaceous, conforming a regional planation surface (Aragón et al., 2010, 2014; Aguilera et al., 2014). Fig. 2A displays a simplified pre-K/P distribution of pre-Jurassic basement, Jurassic volcanic rocks (Chon Aike SLIP), and Cretaceous sedimentary basins in Patagonia.

The first K/P Atlantic transgression covered a significant part of Patagonia (i.e., Feruglio, 1949; Camacho, 1967, 1992; Yrigoyen, 1969; Spalletti et al., 1993; Malumián and Caramés, 1995; Malumián, 1999, 2002; Nández and Malumián, 2008; Aguirre-Urreta et al., 2011; Scasso et al., 2012; Foix et al., 2021); paleogeographic estimations from foraminifera record indicate that about 500.000 km<sup>2</sup> of the current emerged area was flooded by the K/P sea (Fig. 2B), exceeding any other Cenozoic transgression in Patagonia (Malumián and Caramés, 1995).

Most complete marine stratigraphic records took place over Mesozoic sedimentary basins (i.e., Malumián and Caramés, 1995; Malumián, 1999; Nández and Malumián, 2008), but the continental flooding also covered the NPM (Feruglio, 1949; Camacho, 1992; Spalletti et al., 1993; Malumián and Caramés, 1995; Aragón et al., 2010, 2014; Scasso et al.,



**Fig. 2.** A) Simplified geologic configuration before the K/P transgression in Patagonia. Location of the study area. B) Paleogeography of the Maastrichtian transgression (taken from Nández and Malumián, 2008) with flooded and emerged areas. References: 1) Neuquén basin, 2) Salado basin, 3) Northpatagonian Massif, 4) Cañadón Asfalto basin, 5) Valdés basin, 6) Rawson basin, 7) Golfo San Jorge basin, 8) Deseado Region, 9) Austral-Magallanes basin, 10) Malvinas basin, 11) High Andes (Tunik et al., 2004) and 12) Colorado basin. A) Tandilia High (taken from Lovecchio et al., 2018), B) Camarones High, C) San Bernardo Fold-Belt, D) Eastern Patagonian High, E) Deseado High, F) Malvinas High, G) Dungeness High, H) Northern Patagonian Andes and I) Southern Patagonian Andes. Location of the study area on the Northpatagonian Massif.

2012; Foix et al., 2021). Roughly, the K/P marine stratigraphic record in the sedimentary basins is a 200–300 m thick, siliciclastic tidally-dominated succession overlying Cretaceous continental units, mainly characterized by quartz-glaucinitic, transgressive/regressive sandstones (few tens meters thick), inner shelf mudstones (up to 150–200 m thick) and scarce bioclastic deposits (Legarreta et al., 1990; Legarreta and Uliana, 1994; Malumián, 1999; Foix et al., 2021; and cites therein). The K/P marine flooding in the NPM mainly took place over a Middle Jurassic to the Late Cretaceous erosive planation surface (Aragón et al., 2010, 2014; Aguilera et al., 2014). Gianni et al. (2018a, 2018b) suggested that a Late Cretaceous-Paleocene regional subduction-related subsidence (flexural + dynamic) drove this marine flooding over Patagonia during a large flat-slab event.

Several equivalent lithostratigraphic units have been defined for the K/P marine flooding over extra-Andean Patagonia (Feruglio, 1949; Camacho, 1967, 1992; Yrigoyen, 1969; Spalletti et al., 1993; Malumián and Caramés, 1995; Malumián, 1999, 2002; Nández and Malumián, 2008; Scasso et al., 2012; Foix et al., 2021), such as Salamanca Formation (Golfo San Jorge basin), Roca Formation (Neuquén basin-NPM), Arroyo Salado and El Fuerte formations (central-eastern NPM), Bustamante formation (southern NPM), La Colonia formation (Cañadón Asfalto basin), among others. K/P near-shore carbonate deposits over the NPM mainly comprise bioclastic rocks and were included in the “Northpatagonian platform” (Spalletti et al., 1993). Offshore epiclastic deposits usually cover this carbonate record during the transgression deepening

(Feruglio, 1949; Andreis et al., 1975; Spalletti et al., 1993; Ardolino et al., 2003; Foix et al., 2015).

The K/P transgression contains an abundant paleontologic record in the central-northern Patagonia (Table 1). Particularly, the molluscan Danian assemblage in Patagonia incorporates 35 genera of gastropods and 31 genera of bivalves, but Oysters constitute the dominant molluscan group throughout almost all basins (del Río, 2021). Also, these Oyster communities support a wide variety of boring and encrusting biota, including sponges, polychaetes, bivalves, fungi, algae, barnacles, and bryozoans (Brezina et al., 2014, 2017).

Maastrichtian and Danian foraminiferal assemblages were recognized in the marine record (i.e., Kaasschieter, 1963; Méndez, 1966; Chebli and Serraiotto, 1974; Malumián and Caramés, 1995; Nández and Malumián, 2008; Malumián and Nández, 2011; Simeoni, 2014) (see Table 2). Most of Maastrichtian endemic calcareous foraminiferal species disappear during the K/P transition (Malumián and Nández, 2011). Later, Danian cosmopolitan Midway type foraminiferal assemblages of Patagonia include at least 234 benthic endemic species (Malumián and Caramés, 1995).

### 3. Material and methods

We present a regional dataset constituted by both previous works and own information. This contribution includes a regional stratigraphical, lithological, and paleontological characterization of initial transgressive

**Table 1**  
Main K/P marine paleontologic record in central-northern Patagonia.

Fossil record	References
Bivalves	Burckhardt (1901), von Ihering (1903), Celeste (1940), Feruglio (1936, 1949), Petersen (1946), Camacho (1992), Spalletti et al. (1993), Griffin and Hünicken (1994), Casadio (1998), del Río (2004, 2012, 2021), Griffin et al. (2005, 2008), del Río et al. (2011), Scasso et al. (2012), del Río and Martínez (2015)
Gastropods	Feruglio (1949), Camacho (1992), Griffin and Hünicken (1994), del Río (2012)
Brachiopods	Feruglio (1949), Chebli and Serraiotto (1974), Spalletti et al. (1993)
Echinoderms	Feruglio (1949), Parma (1989), Spalletti et al. (1993) Parma and Casadio (2005), Martínez et al. (2011), Foix et al. (2015)
Barnacles	Brezina et al. (2014, 2017)
Briozoans	Brezina et al. (2014, 2021)
Coral reefs and stromatolites	Baron-Szabo et al. (2003), Kiessling et al. (2005), Aguirre-Urreta et al. (2011), Scasso et al. (2012), Carrera and Casadio (2016)
Boring polychaetes	Brezina et al. (2014)
Nannofossils, foraminifera, palynomorphs and dinoflagellates	Méndez (1966), Bertels (1969, 1973, 1974, 1975, 1979, 1995), Archangelsky (1973, 1976), Chebli and Serraiotto (1974), Malumián et al. (1983), Concheyro and Nández (1994), Malumián (1999, 2002), Palamarczuk et al. (2002), Nández and Malumián (2008), Malumián and Nández (2011), del Río et al. (2011), Simeoni (2014), Vellekoop et al. (2017), Guler et al. (2014, 2019)

K/P marine sedimentary rocks and their substrate in the NPM. Freely accessible satellite images were used to characterize some stratigraphic relationships. The regional cartography of the studied rocks was made using a free Geographic Information System (QGIS). Complementary, thin sections were used for qualitative petrographic characterizations.

#### 4. Results

We analyzed the basal K/P transgressive stratigraphic record over 70,000 km<sup>2</sup> in the NPM, central-northern extra-Andean Patagonia (Argentina). 64 localities were described/compiled in the Chubut and Río Negro provinces (Fig. 3), mainly conformed by isolated outcrops. The substrate of K/P marine deposits is constituted by Jurassic volcanic rocks (83 % = 53 localities), Cambrian metamorphic rocks (12 % = 8 localities), Triassic volcanic rocks (3 % = 2 localities) and Ordovician plutonic rocks (2 % = 1 locality).

##### 4.1. Chubut province

In the south-eastern extreme of the NPM (Fig. 3), near the Atlantic seacoast, the K/P bioclastic record was named Rocanense by Feruglio (1936, 1949), Bustamante Formation (Simpson, 1941; Ardolino et al., 2003) or Bustamante Member/Salamanca Fm (Andreis et al., 1975; Scitutto et al., 2000; Foix et al., 2015). We described this bioclastic unit in two carbonate quarries: Cantera El Tablón (locality 4) and Tetas de Pineda (locality 1). At El Tablón quarry, the bioclastic deposits overlie the MVC (Fig. 4A, B), even preserving mollusc shells (oysters?) on the unconformity (Fig. 4C, D). In this locality, the carbonate record reaches

up to 10 m in thickness (Ardolino et al., 2003) being characterized by bioclasts ranging between 0.1 and 10 cm in large (Fig. 5A) with some volcanic clasts (1–4 cm) near the basal contact (Fig. 5B). The bioclastic deposits display meter-scale low-angle cross-bedding and reactivation surfaces dipping away from the Jurassic outcrops, interpreted as fore-shore deposits over paleo-islands (Fig. 5C). Bivalves with minor bryozoan and echinoderms are the most abundant marine fossil remains in the bioclastic deposits (Fig. 6). Overlying the basal bioclastic interval, the stratigraphic succession continues with inner shelf mudstones deposited during the marine deepening.

At Tetas de Pineda (locality 1), the Salamanca Formation also overlies the MVC with a 15–20 m thick, parallel stratified, basal bioclastic stratigraphic interval that surrounds the local paleohighs (Fig. 7A, B). Oyster remains are frequent and some volcanic clasts up to 3–4 cm in diameter were described (Fig. 7C). Petrographic analysis of thin sections shows skeletal packstones mainly constituted by bivalve remains with felsitic volcanic lithoclasts (Fig. 8A, B). Also, echinoderm plates (Fig. 8C) and spines (Fig. 8D, E), coralline red algae (Fig. 8D), bryozoan (Fig. 8E) and benthic foraminifera (Fig. 8F) were recognized.

Other nearby quarries, such as Bahía Bustamante (locality 5) and La Esther (locality 7), contain bioclastic deposits about 4 m and 9 m in thickness respectively (Ardolino et al., 2003). The stratigraphic relationship between MVC/bioclastic deposits was also described at the Aristizábal (locality 2) and Gravina (locality 3) peninsulas (Feruglio, 1949; Ardolino et al., 2003), with bioclastic deposits ranging 4–5 m in thickness (Ardolino et al., 2003).

On the banks of both lower Chico and Chubut rivers (Fig. 3), the K/P bioclastic deposits also overlie Jurassic volcanic rocks (Feruglio, 1949;

**Table 2**  
K/P foraminifera record in central-northern Patagonia.

Age	Benthic foraminifera	Planktonic foraminifera	References
Maastricht.	<i>Discorbis correcta</i> <i>Bulliminella isabelleana</i> <i>Charltonina acutimarginata</i> <i>Angulogavelinella?</i> sp.		Nández and Malumián (2008)
Danian	<i>Bulliminella isabelleana</i> <i>Lagenoglandulina neuquensis</i> <i>Favolagena atilai</i> <i>Ammoelphiidiella?</i> sp. <i>Discorbinella castellaroeae</i> <i>Guttulina luisae</i> <i>Lagena archangelsky</i> <i>Lenticulina wichmanni</i> <i>Migros hanseni</i> <i>Palmula budensis rocanense</i>	<i>Globigerina pseudobulloides</i> <i>Globigerina triloculinoidea</i> <i>Globigerina compressa</i> <i>Globigerina daubjergensis</i> <i>Chiloguembelina midwayensis</i> <i>Globoconusa daubjergensis</i>	Kaasschieter (1963) Chebli and Serraiotto (1974)

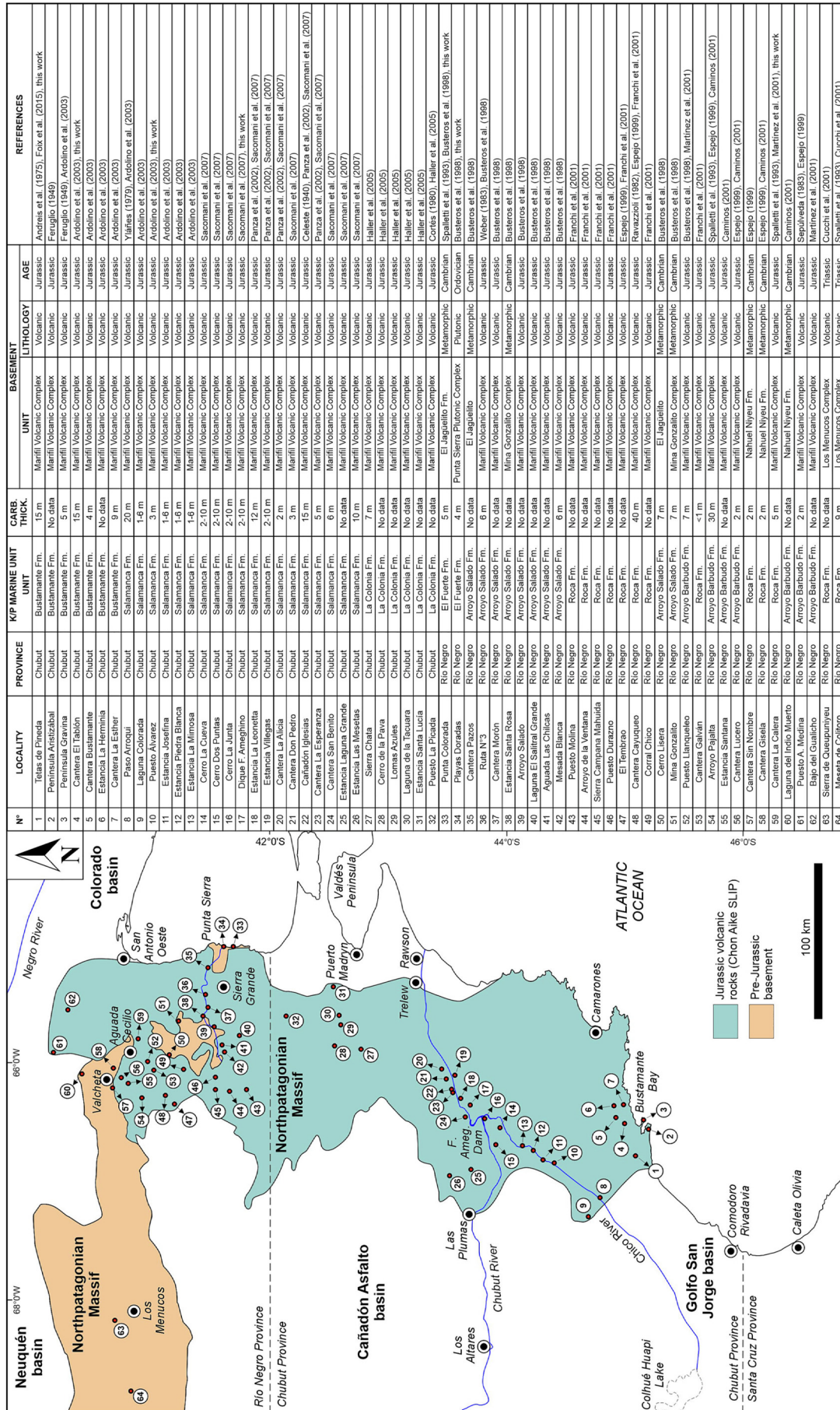
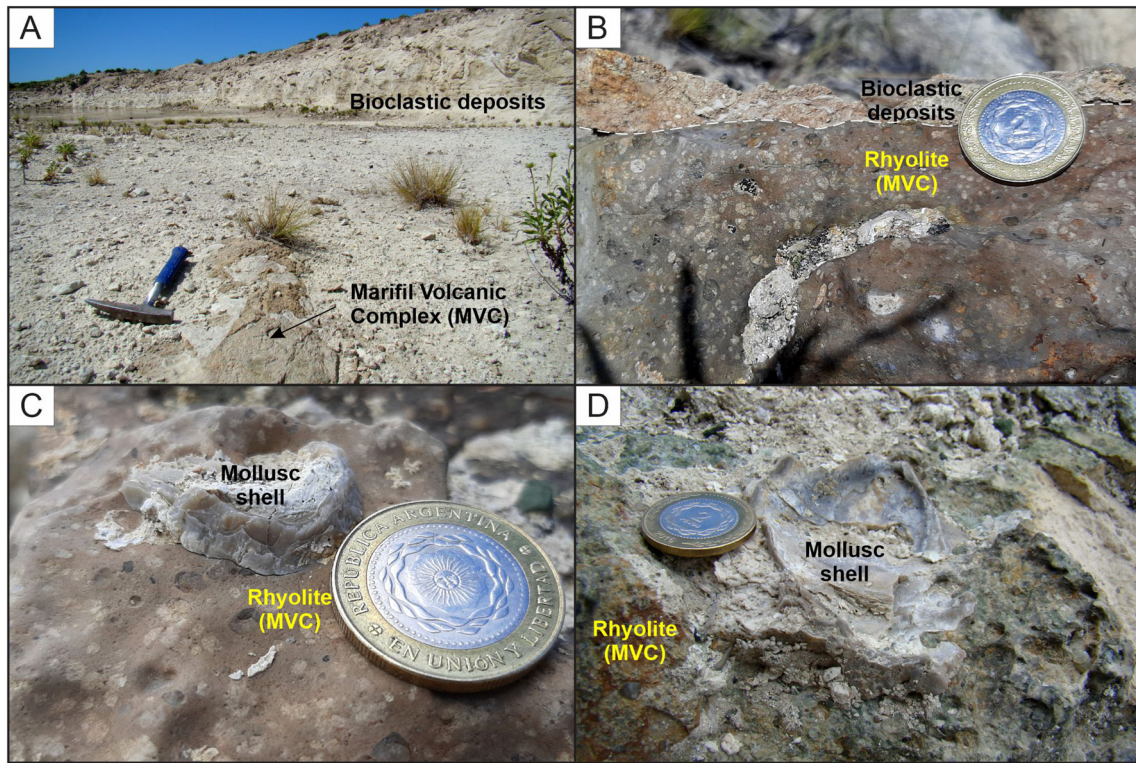


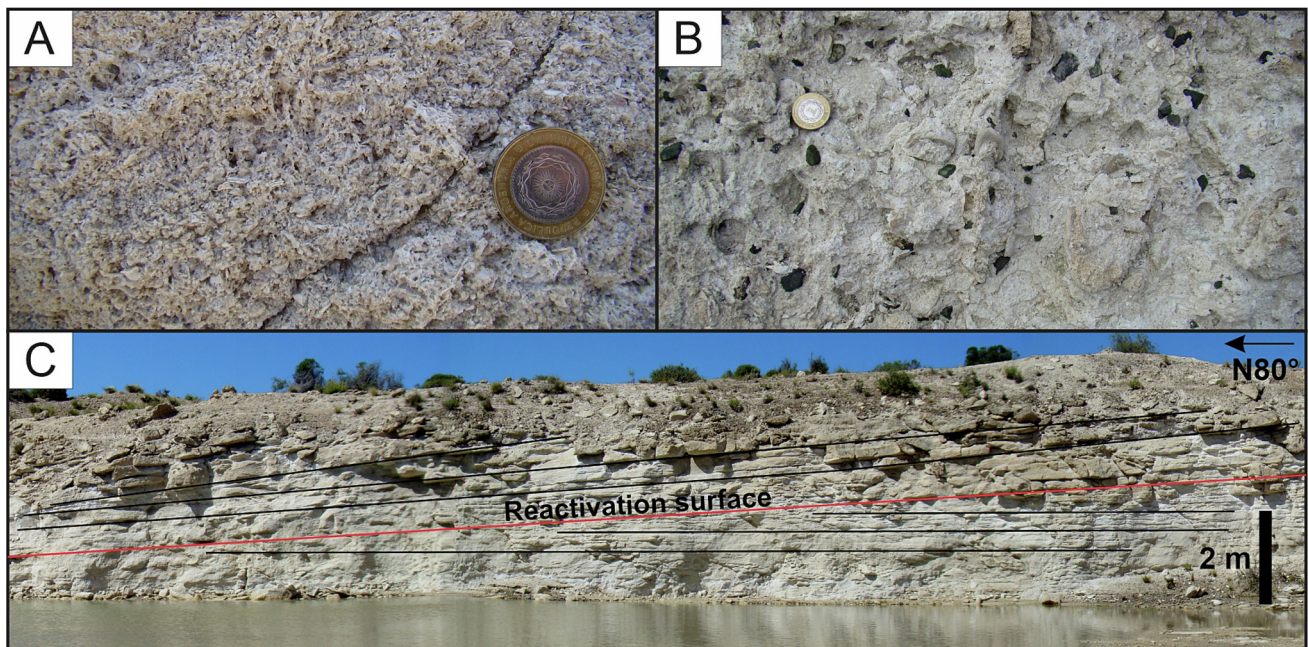
Fig. 3. Study area and distribution of localities where the K/P marine record overlies Jurassic (Chon Aike SLIP) or pre-Jurassic basement in the NPM.



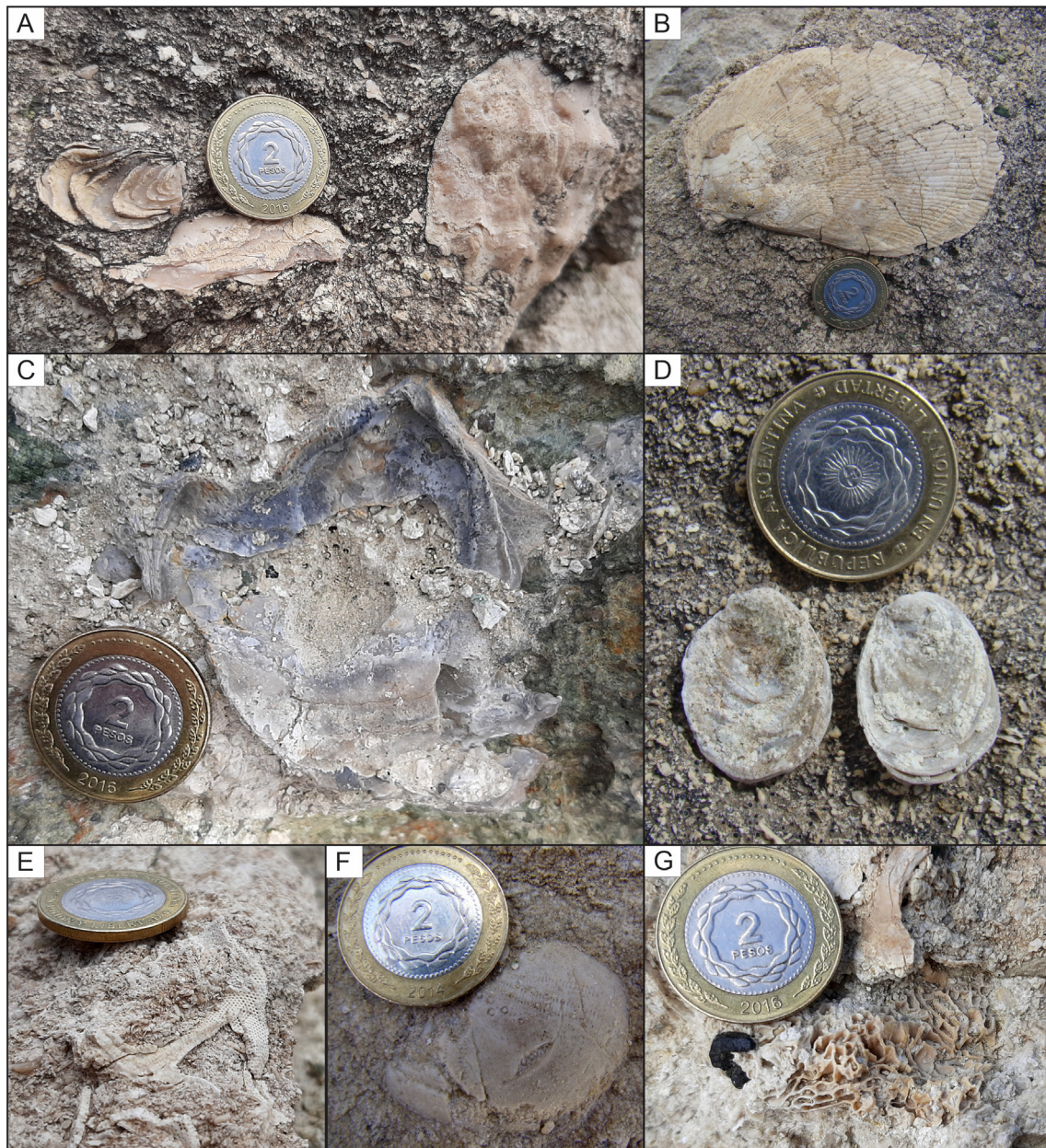
**Fig. 4.** Unconformity between Jurassic volcanic rocks (VMC) and K/P bioclastic deposits, El Tablón quarry (Chubut). A, B) The bioclastic carbonates over the Jurassic rhyolitic rocks (VMC). C, D) Mollusc shells (Oysters) over the Jurassic volcanic rocks (VMC). The coin is 2.3 cm in diameter.

Ylláñez, 1979; Panza et al., 2002; Ardolino et al., 2003; Sacomani et al., 2007). At the Paso Arroqui (locality 8), bioclastic deposits up to 20 m thick are interbedded with sandstone (Ylláñez, 1979 in Sacomani et al., 2007). At the Puesto Álvarez (locality 10), the basal sedimentary record is composed of a 3 m thick bioclastic sandstones with large-scale cross-stratification (Fig. 9A). In this case, the substrate (MVC)

seems to preserve a 0.5 m thick altered profile with volcanic boulders (Fig. 9B). In the lower Chubut River valley, downstream of the Florentino Ameghino dam, several carbonate quarries have been exploited with carbonate thickness ranging 2–6 m (Panza et al., 2002; Sacomani et al., 2007): Cantera La Alicia (locality 20), Cantera Don Pedro (locality 21), Cantera La Esperanza (locality 23), Cantera San



**Fig. 5.** Basal bioclastic deposits at El Tablón quarry. A) Very fragmented bioclastic carbonates. B) Frequent volcanic clasts near the basal contact with de MVC. C) Meter-scale low-angle cross-bedding and reactivation surface interpreted as foreshore deposits. The coin is 2.3 cm in diameter.



**Fig. 6.** Invertebrate fauna in K/P bioclastic deposits (El Tablón quarry). A–D) Bivalve remains, including Oysters. E) Bryozoan. F) Echinoderms (*Linthia*?). G) Corals? remains. The coin is 2.3 cm in diameter.

Benito (locality 24). At the Cañadón Iglesias outcrops (Feruglio, 1949; Lapido and Page, 1979; Lapido, 1981), Celeste (1940) described a bioclastic record up to 15 m in thickness (locality 22).

Tabular outcrops of the K/P bioclastic deposits often conform structural terraces over the MVC, such as at Dique Florentino Ameghino (locality 17, Fig. 10A) or near Estancia Las Mesetas (locality 26, Fig. 10B) with 10 m in thickness (Sacamani et al., 2007).

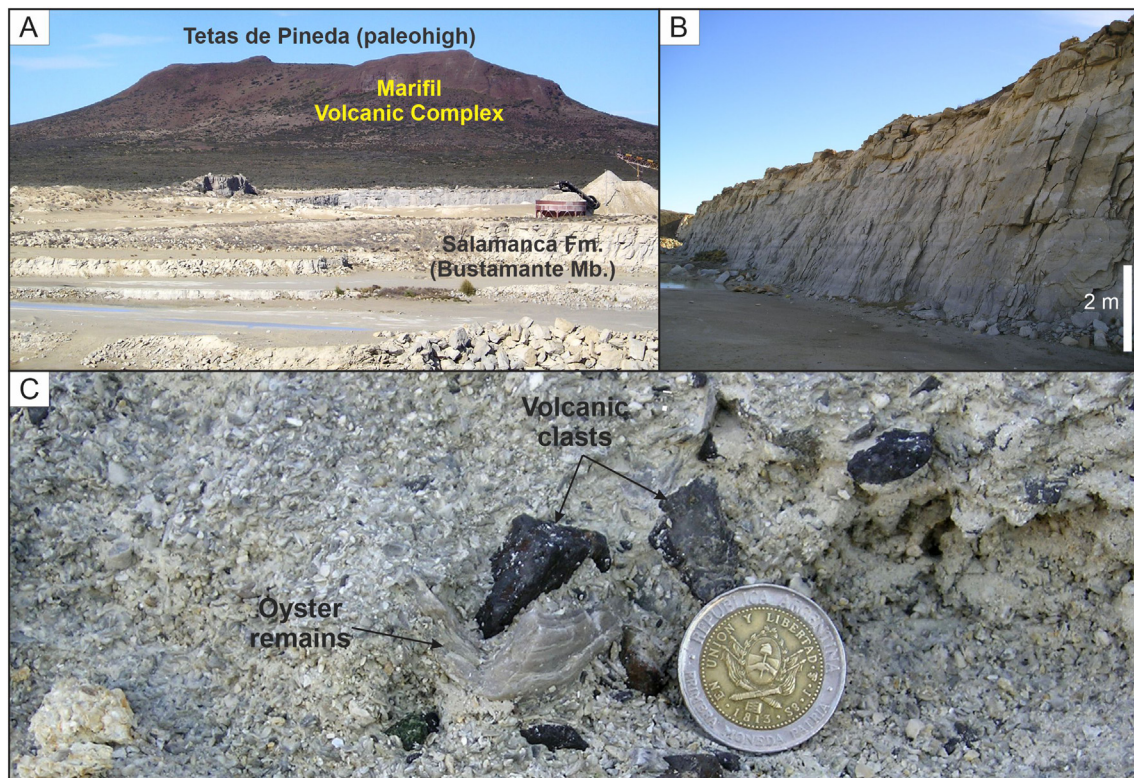
Westward of Puerto Madryn city (Fig. 3), the K/P marine deposits were described as Puesto La Picada (Cortés, 1980) or La Colonia Formation (Haller, 1981; Haller et al., 2005) (locality 32). Near Sierra Chata (locality 27), Haller et al. (2005) described the unconformity between the marine sedimentary and the Jurassic VMC, where the transgressive record includes mudstones and silicified carbonates up to 7 m thick.

Feruglio (1949) resumed the marine paleontological record of this carbonate interval over the Chubut province, including bivalves (i.e., *Pectunculus feruglioi*, *Trigonia wilckensi*, *Venericardia*

*palaeopatagonica*, *Patagocardia peterseni*, *Phacoides* sp., *Lucina* sp., *Meretrix rothi*, *Solecurtus* sp., *Panopea* sp., *Lima* sp., *Pecten* sp., *Ostrea ameghinoi*, *Gryphaea rostrigera*, *Exogyra mendozana*, *Modiola aprilis*), schaphopods (i.e., *Dentalium* sp.), gastropods (i.e., *Natica* sp., *Ampullospira dubia*, *Calyptreaea* sp., *Scalaria* sp., *Turritella malaspinga*, *Melania ameghiniana*, *Arrhoges gregaria*, *Perissoptera* sp., *Triton* sp., *Cominella praecursor*, *Retusa sculata*, *Cinulia pauper*) and echinoids (i.e., *Linthia joannis-böhmi*).

#### 4.2. Río Negro province

In the northeastern NPM, the K/P transgression conformed a widely distributed carbonate platform composed of skeletal fragments (Spalletti et al., 1993), including Roca, Arroyo Barbudo, El Fuerte and Arroyo Salado formations (Fig. 3). Along the Salado creek, northward of Sierra Grande city, the Arroyo Salado Formation shows several carbonate outcrops overlying both Paleozoic igneous-metamorphic



**Fig. 7.** Basal bioclastic deposits of the Bustamante Formation/Member (Salamanca Fm.) over the MVC (Tetas de Pineda quarry), Chubut (modified from Foix et al., 2015). A) Tetas de Pineda quarry. B) Tabular strata of bioclastic carbonates. C) Domain of bioclasts of bivalve molluscs (Oysters?) with the presence of volcanic clasts (substrate). The coin is 2.3 cm in diameter.

basement and the MVC (Busteros et al., 1998) (Fig. 3). The unconformity over Cambrian high-grade metamorphic rocks (locality 38) and Jurassic volcanic rocks (locality 32) can be observed on satellite images (Fig. 11A, B).

Nearly Valcheta city (Fig. 3), the Arroyo Barbudo/Roca Formation usually overlies metamorphic rocks of the Nahuel Niyeu Formation (localities 57, 58 and 60) or the MVC (localities 53, 56 and 59), often used as carbonate exploitations (Espejo, 1999; Caminos, 2001). The Cantera La Calera (locality 59) is a carbonate quarry that contains massive, silicified carbonates up to 5 m thick (Fig. 12). The Arroyo Barbudo Formation contains a bioclastic basal record up to 30 m thick over the MVC (Espejo, 1999; Caminos, 2001) at the upper Pajalta Creek (locality 54), with equivalents in the Gran Bajo del Gualicho (localities 61 and 62) (Sepúlveda, 1983; Martínez et al., 2001). This relationship also occurs in isolated outcrops of the Roca/Arroyo Barbudo Formation surrounding the Somouncurá basalts, such as Ventana creek (locality 44), Sierra Campana Mahuida (locality 45), Puesto Durazo (locality 46), El Tembrao (locality 47), Corral Chico (locality 49) (Franchi et al., 2001; Caminos, 2001). At Cantera Cayuqueo (locality 48), Ravazzoli (1982) describes 40 m of carbonate sandstones overlying the MVC.

The El Fuerte Formation outcrops on the Atlantic seacoast between San Antonio Oeste city and Punta Colorada with carbonate sandstones ranging about 4–6 m in thickness (Rodríguez, 1990; Busteros et al., 1998), but reaching up to 9 m few kilometers northward of Punta Sierra (Weber, 1983). At the mouth of Salado Creek and Playas Doradas (locality 34), the El Fuerte Formation overlies El Salado pluton (Fig. 13A, B), an intrusive body that integrates the Ordovician Punta Sierra Plutonic Complex (Busteros et al., 1998). At the Punta Colorada (locality 33) is outcropped the unconformity between low-angle cross-bedded bioclastic sandstones (El Fuerte Formation) and the Cambrian, low-grade metamorphic rocks of the El Jagüelito Formation (Fig. 13C).

In the central-western MNP, the Roca Formation overlies volcanic/volcaniclastic rocks of the Los Menucos Complex (Fig. 3) and contains

up to 9 m of bioclastic deposits (Spalletti et al., 1993; Cucchi et al., 2001) at the Meseta de Colitoro (locality 64).

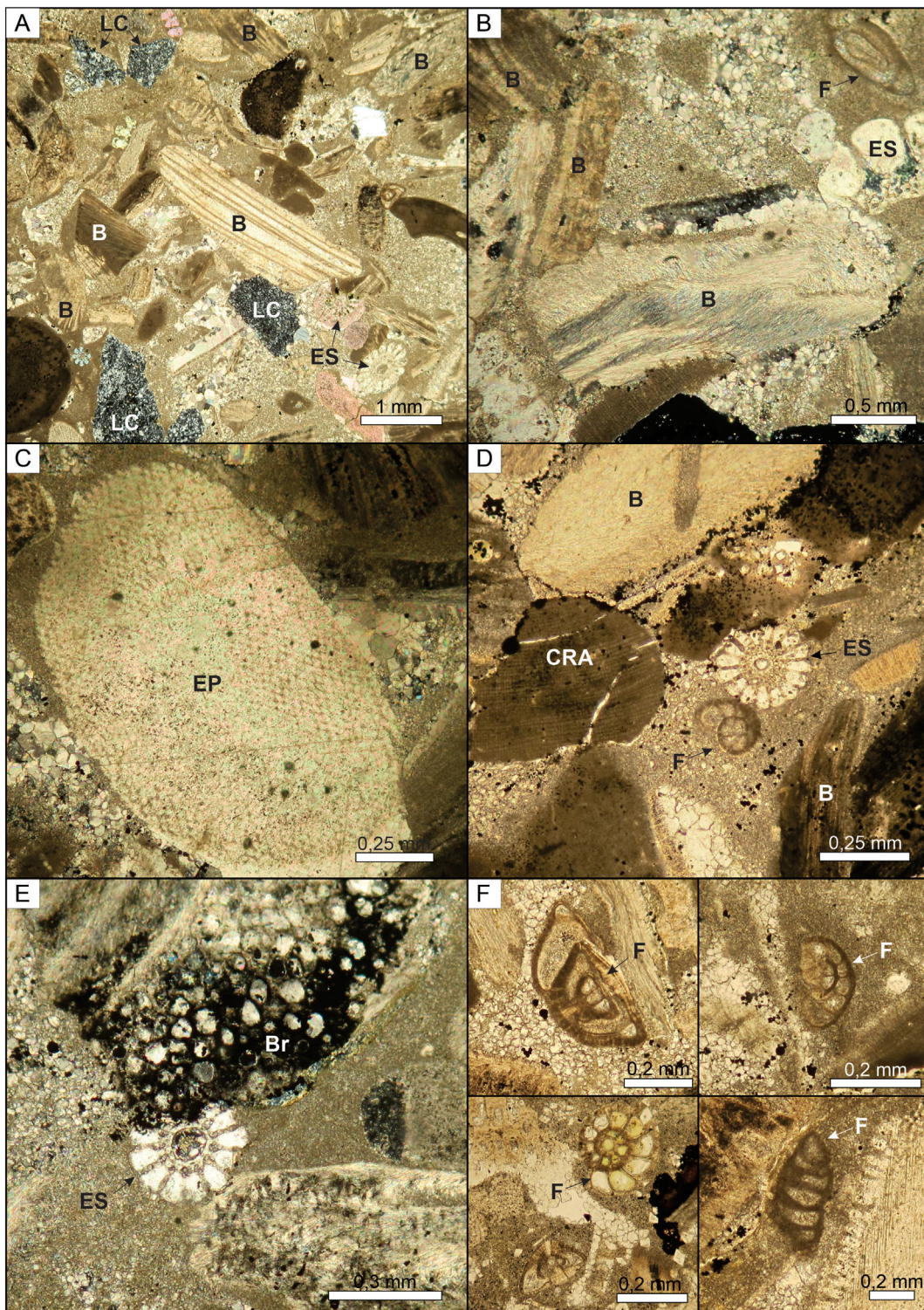
In this region of the NPM, Spalletti et al. (1993) described a variable paleontological record, including *Gryphaea* sp., *Gryphaea rothi*, *Odontogryphaea*, *Exogyra* sp., *Ostrea clarae*, *Ostrea rionegrensis*, *Venericardia* sp., *Leda*, pectinids and terebratulid brachiopods.

#### 4.3. Summary and interpretations

The work includes 64 localities where the K/P bioclastic marine record overlies volcanic (MVC - Chon Aike SLIP) or igneous-metamorphic rocks over ~70,000 km<sup>2</sup>. These results show a regional rocky shore occurred during a major transgressive event over rocky reliefs, the right paleoenvironmental scenario for their formation (i.e., Surlyk and Christensen, 1974; Bromley and Asgaard, 1993; Webb, 1994; Jia-yu and Johnson, 1996; Sanders, 1998; de Gibert et al., 1998, 2012; Betzler et al., 2000; Larsen et al., 2003; D'Alessandro et al., 2004; Shepard, 2006; Felton et al., 2006; Surlyk and Sørensen, 2010; Buatois and Encinas, 2011; Bover-Arnal et al., 2011; Brlek et al., 2018; Sanders et al., 2019). This extra-Andean ancient rocky shore is the widest known in South America, and here is named as “Northpatagonian rocky shore” (Fig. 14).

The sedimentary record of rocky shores usually contains bioclastic deposits because they constitute a favorable environment for benthic carbonate production (cf. Surlyk and Christensen, 1974; Surlyk, 1997; Libbey and Johnson, 1997; de Gibert et al., 1998, 2012; Sanders, 1998; Betzler et al., 2000; Cebrián et al., 2000; Johnson and Baarli, 1999, 2012; Taylor and Wilson, 2003; D'Alessandro et al., 2004; Surlyk and Sørensen, 2010; Sørensen and Surlyk, 2013, 2015; Sanders et al., 2019), as we have described in the K/P studied example. We suppose that the molluscan death assemblage has a high environmental fidelity to the rocky shore life assemblage, but the low representation of gastropods can also result from differential taphonomic loss (Sørensen and Surlyk, 2015).

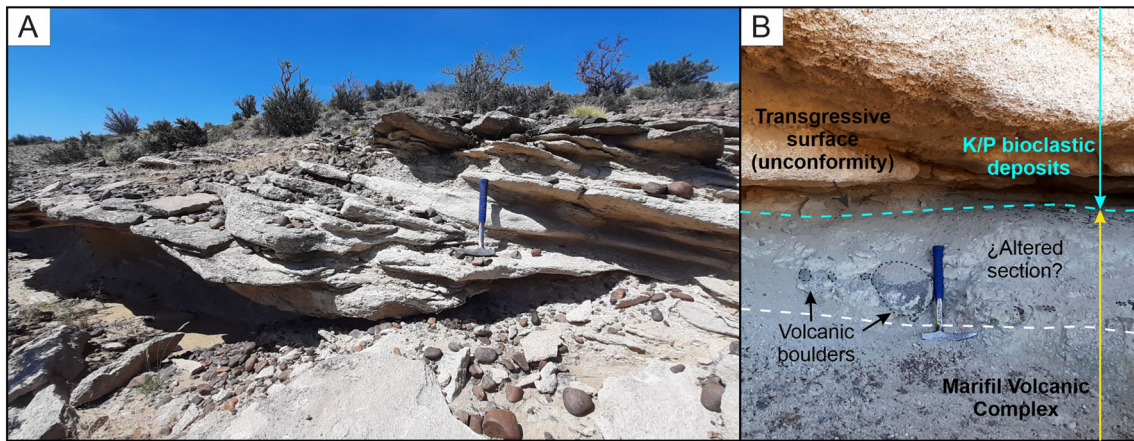




**Fig. 8.** Thin-section photomicrographs of bioclastic deposits (locality 1, Tetas de Pineda). A) Skeletal packstone mainly constituted by bivalve remains (B), with presence of felsitic volcanic lithoclasts (LC). Sparry calcite cement. Crossed nicols. B) Foliated structure of bivalve shell (B), benthic foraminifera (F) and fragment of echinoderm spine (ES). Sparry calcite cement. Crossed nicols. C) Fragment of echinoderm plate (EP) with uniform “honeycomb” microtexture. Crossed nicols. D) Coralline red algal grain (CRA) showing a fine-scale reticulate structure. Transverse section of an echinoderm spine with single-crystal optical behavior. Parallel nicols. E) Tangential section of bryozoan fragment (Br) with a regular box-like arrangement (note sparry carbonate infilling the porous). A transverse section of echinoderm spine is also observed. Crossed nicols. F) Multicamerate benthic foraminifera (F). Parallel nicols.

Most of encrusting communities are restricted to shallow marine environments, colonizing subtidal and intertidal zones (cf. Bromley and Asgaard, 1993; Libbey and Johnson, 1997; Cruz-Motta et al., 2010; Santos et al., 2011; Bover-Arnal et al., 2011; Brlek et al., 2018). In this way, the dominant macrobiota of K/P bioclastic deposits allows us to

interpret intertidal and shallow subtidal paleoenvironmental conditions during the basement flooding over a very wide area of the extra-Andean Patagonia. Considering a tabular thickness of about 2–4 m covering only half of the Northpatagonian rocky shore (35.000 km<sup>2</sup>), the volume of K/P bioclastic carbonate deposits could be about 70–140 km<sup>3</sup>.



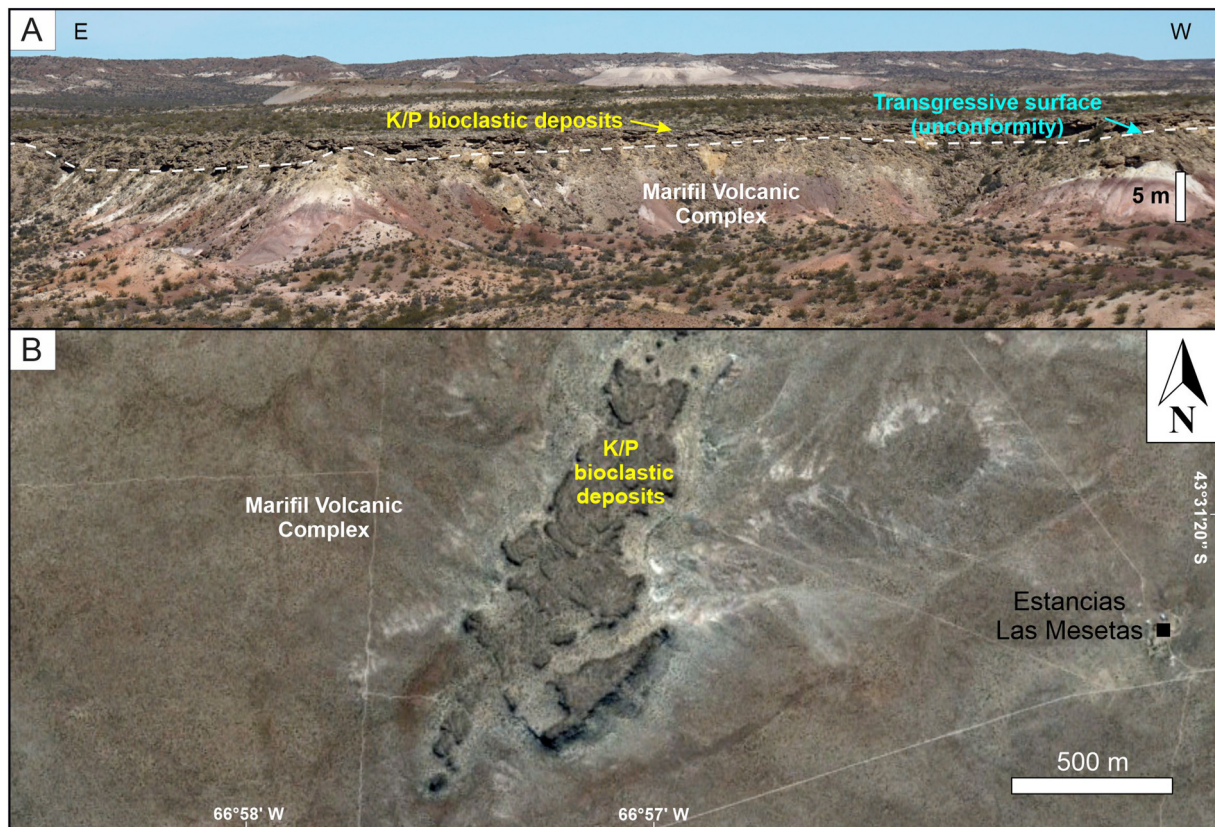
**Fig. 9.** A) Large-scale cross-bedded bioclastic sandstones, Puesto Álvarez locality (Chico River). B) Transgressive surface over acid volcanic rocks (MVC), which seems to preserve a 0.5 m thick altered profile, Puesto Álvarez locality (Chico River). B).

**5. Discussion**

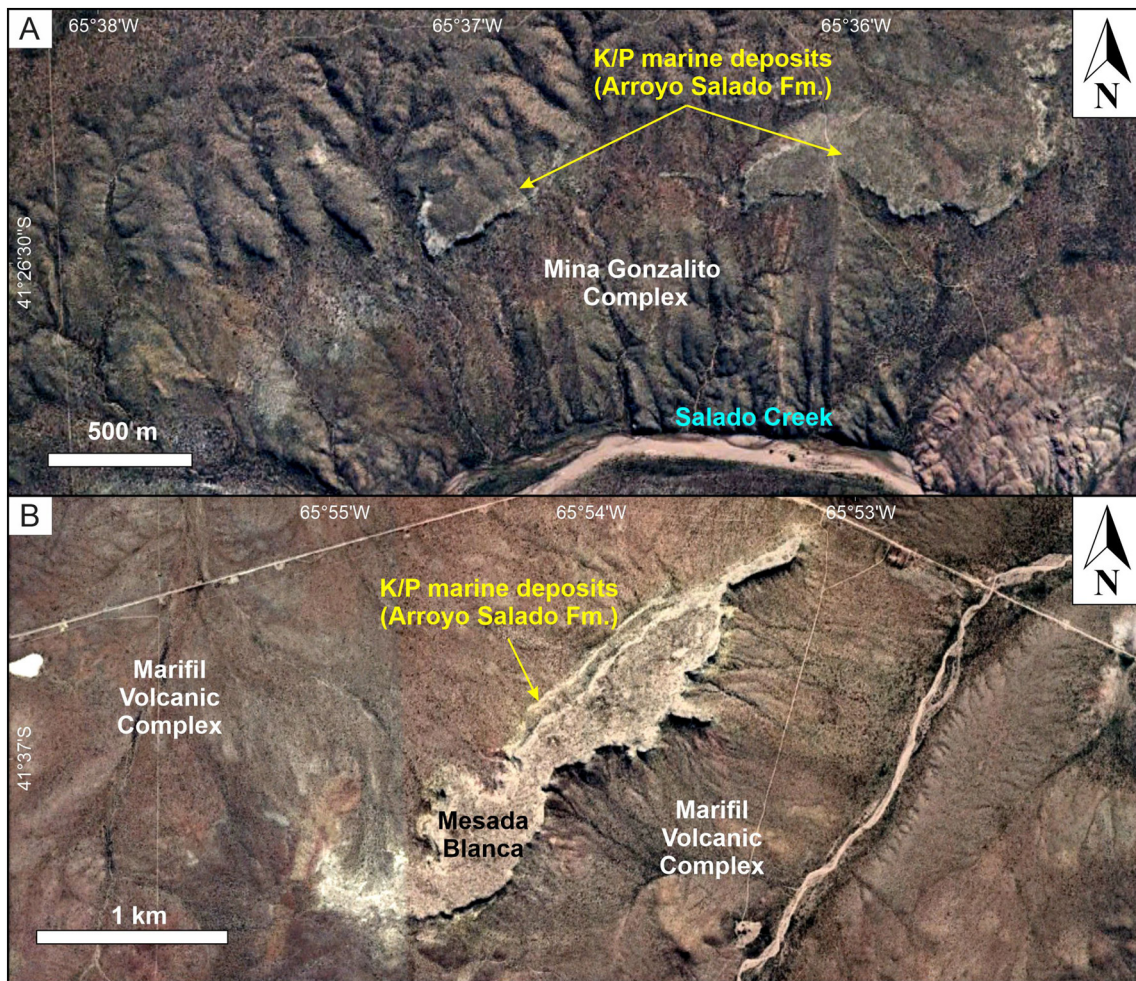
Long-term continental flooding episodes are mainly controlled by sea-level rise (i.e., Hays and Pitman, 1973; Vail et al., 1977; Haq et al., 1987, 1988), regional subsidence (i.e., Gurnis, 1991, 1993; Müller et al., 2008, 2018) or a combination of both (Miller et al., 2005; Spasojevic and Gurnis, 2012; Haq, 2014; Dávila et al., 2018, 2019; Cao et al., 2019). Highest frequency of rocky shore occurs on volcanic islands, active orogenic wedges and convergent-plate margins (Johnson, 1988a), with very well documented examples (Johnson et al., 1996; Sanders, 1998; Felton, 2002; Spencer and Viles, 2002; D'Alessandro et al., 2004; Felton et al., 2006; Andriani and Walsh,

2007; Bover-Arnal et al., 2011; Buatois and Encinas, 2011; de Gibert et al., 2012; Brlek et al., 2018; Sanders et al., 2019). However, epicontinental flooding usually covers both sedimentary basins and basement rocks, where the basement-onlapping shallow marine successions constitute rocky shores (i.e., Surlyk and Christensen, 1974; Johnson, 1988a, 1988b, 1992; Johnson et al., 1996; Surlyk, 1997; Johnson and Baarli, 1999, 2012; Domènech et al., 2001; Desrochers, 2006; Surlyk and Sørensen, 2010; Sørensen and Surlyk, 2010, 2011, 2013, 2015; Sørensen et al., 2011, 2012; Puig López et al., 2023).

The K/P transgressive marine record over central-northern extra-Andean Patagonia has been studied for over a century. However, most contributions are local studies or monodisciplinary methodological



**Fig. 10.** A) K/P bioclastic deposits overlying the MVC (unconformity), Lower Chubut River (downstream of Florentino Ameghino Dam). Picture courtesy of Juan Manuel Turra. B) Tabular relict outcrops of K/P marine deposits over the MVC, Estancia Las Mesetas (Chubut).

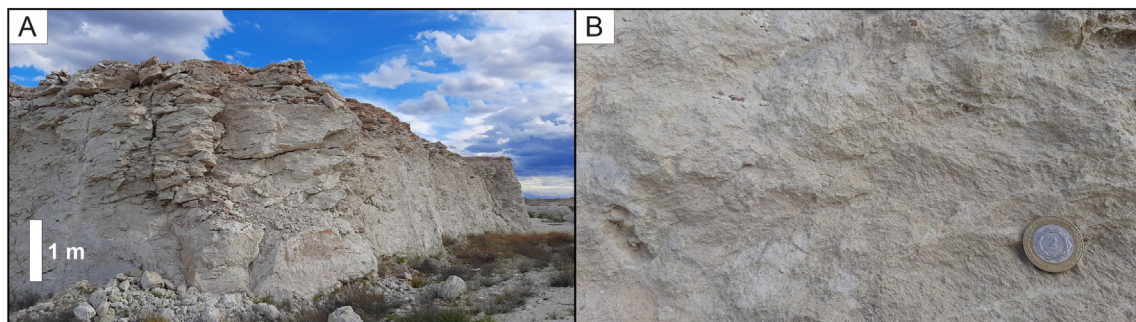


**Fig. 11.** Google Earth satellite images of K/P marine outcrops (Arroyo Salado Fm.) in the NPM, Río Negro Province. A) Unconformity over Paleozoic high-grade metamorphic rocks (Mina Gonzalito Complex), nearly Estancia Santa Rosa. B) Unconformity over Jurassic volcanic rocks (Chon Aike SLIP), Mesada Blanca locality.

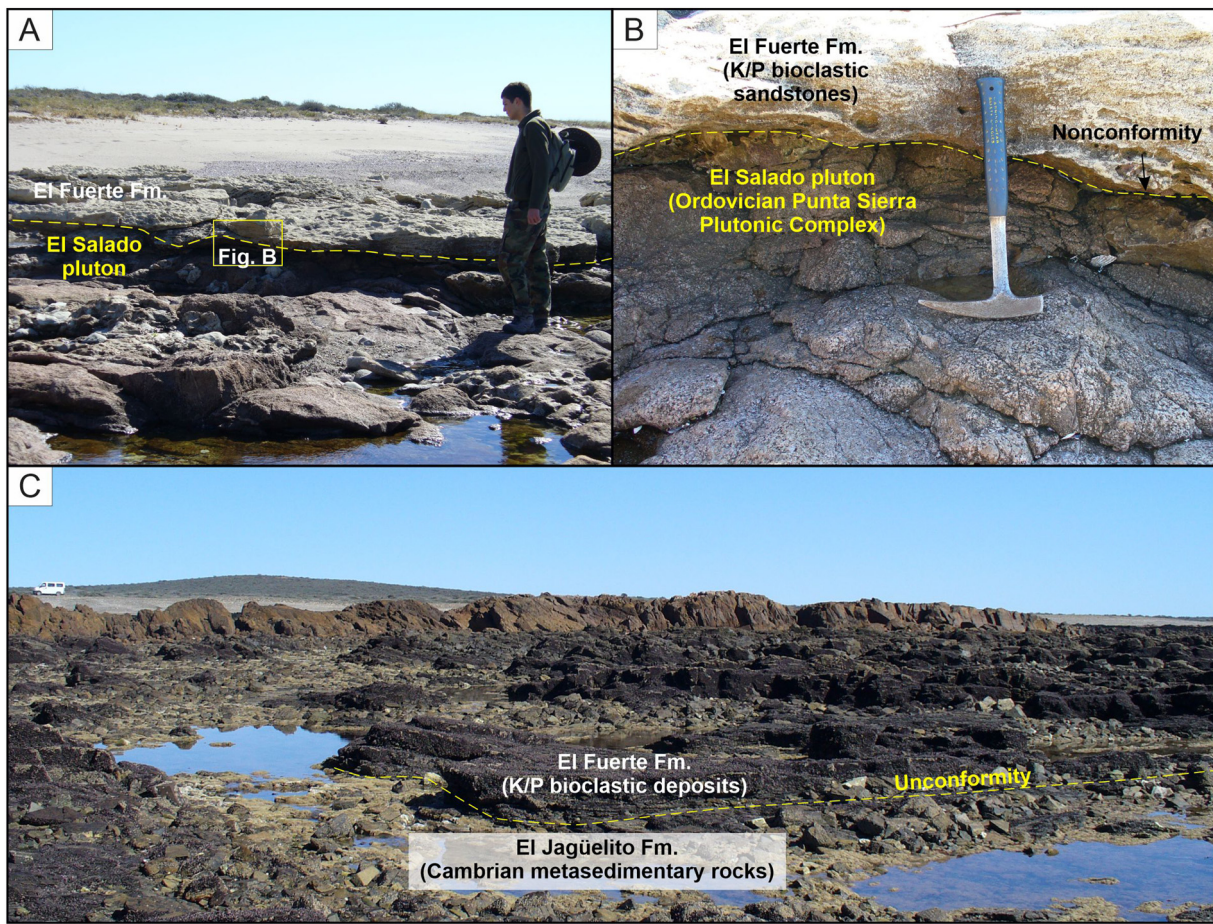
approaches. Although this K/P marine record disposed over the Jurassic and Paleozoic basement probably conforms to the widest outcropped unconformity/nonconformity has not yet been interpreted as a regional rocky shore. This local scenario reinforces the global idea that rocky shores have received insufficient attention from stratigraphers (Johnson, 1988a, 1988b, 1992) and their deposits might be more common in the geological record than previously thought (Puig López et al., 2023). We discussed the main control factors for the K/P Northpatagonian rocky shore, including geological, paleogeographical, and paleoecological implications (Fig. 15).

### 5.1. Geology

The pre-K/P extra-Andean Patagonian geology is a first-order factor on the regional substrate distribution during the continental flooding (Fig. 14). The Chon Aike SLIP constituted the widest substrate for the K/P Northpatagonian rocky shore, with more restricted Paleozoic igneous-metamorphic examples. The basement rocks were flooded by the K/P sea where there was no previous Cretaceous sedimentary record, or where it was eroded (Aragón et al., 2010, 2014; Aguilera et al., 2014). The rapid sea level rise during the advance of



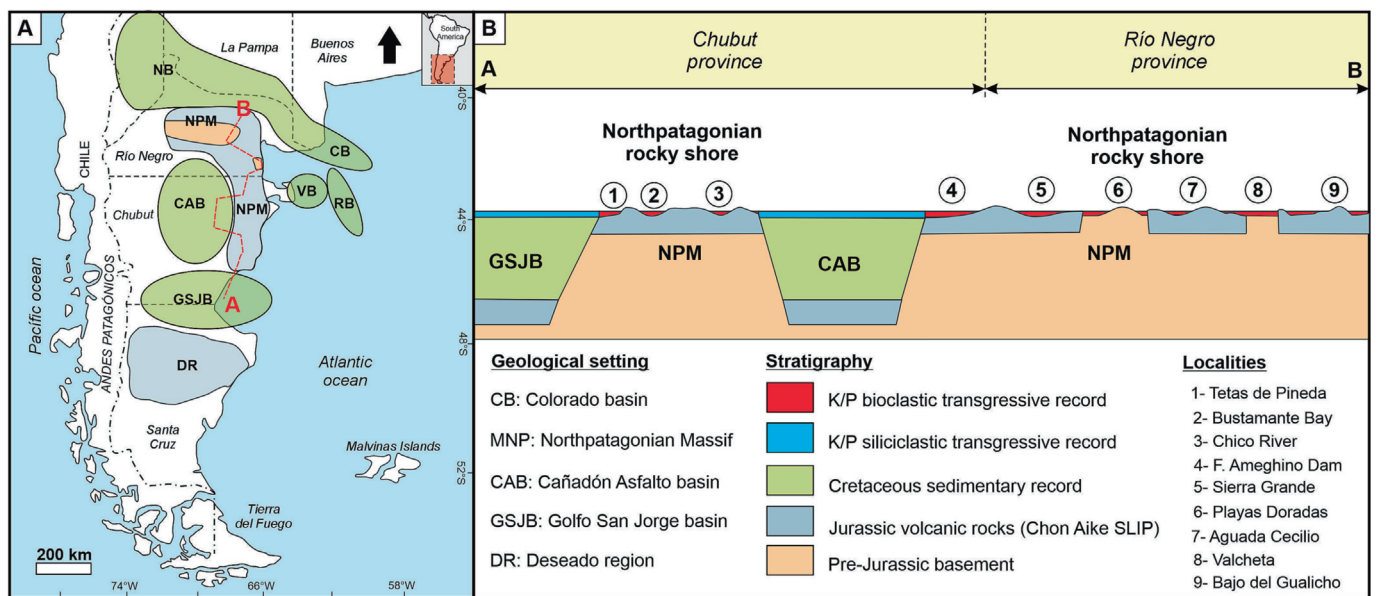
**Fig. 12.** A) The Arroyo Barbudo/Roca Formation at Cantera La Calera, nearly Aguada Cecilio locality. B) Detail of massive, silicified carbonates. The coin is 2.3 cm in diameter.



**Fig. 13.** A) Nonconformity between El Salado pluton (Ordovician Punta Sierra Plutonic Complex) and El Fuerte Formation (K/P bioclastic sandstones), Playas Doradas locality. Location of panel B. B) Detail of the nonconformity between Ordovician igneous basement and bioclastic marine deposits. C) Unconformity between Cambrian, low-grade metamorphic rocks (El Jagüelito Fm.) and the El Fuerte Formation, Punta Colorada locality.

the transgression probably inhibited the development of more encrusting invertebrates, reducing the biodiversity and community net production (i.e., Rilov et al., 2021) and favoring the preservation

of rocky shores (Manikam et al., 2022; Puig López et al., 2023). In addition to providing less favorable substrates to encrusting biota, sedimentary basins (Golfo San Jorge, Cañadón Asfalto, Valdés-Rawson,



**Fig. 14.** A) Simplified distribution of main extra-Andean Mesozoic geological settings, Patagonia (Argentina). Location of A and B sections. B) Regional distribution of K/P transgressive deposits over sedimentary basins and the Northpatagonian Massif (A–B section is about 1000 km in length, without vertical scale). The Northpatagonian rocky shore included at least 70.000 km<sup>2</sup> in the Northpatagonian Massif.

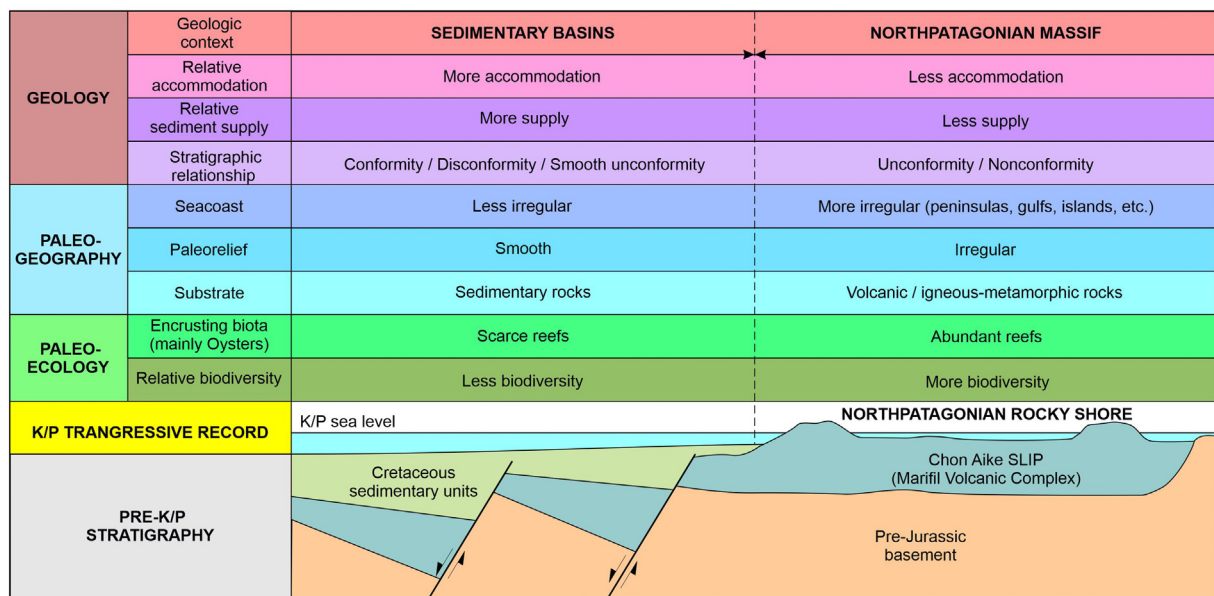


Fig. 15. Main control factors on the K/P transgressive deposits over Mesozoic sedimentary basins and the Northpatagonian Massif, Patagonia (Argentina).

Colorado, Neuquén) had relatively high subsidence rates and sediment supply than the NPM (Fig. 15).

### 5.2. Paleogeography

The type of substrate partially determines the type of biota that may occur. For example, suspension feeders prefer a hard substrate (i.e., Bromley and Asgaard, 1993; de Gibert et al., 2012; Bayne, 2017; Reijmer, 2021). The NPM, and particularly the Chon Aike SLIP, provided a continental-scale hard substrate for the conformation of K/P encrusting rocky shore biota. The pre-K/P relief included a regional erosive planation surface with large, rounded landforms (whaleback type) and isolated inselbergs (Aguilera et al., 2014). The initial K/P marine inundation probably turned the inselbergs into islands, the negative landforms into gulfs/bays and the positive rounded landforms into peninsulas/barriers. Thus, the irregular seacoast configuration (Feruglio, 1949; Foix et al., 2015) would have increased/controlled the ecological niches for encrusting invertebrates (cf. Betzler et al., 2000) and related bioclastic deposits (Fig. 15). A K/P archipelago of volcanic rocks is a very probable paleogeographic scenario for the NPM during the initial transgression because positive landforms have no bioclastic record (i.e., Tetras de Pineda). Though geographic distribution suggests that the world's rocky shores are segregated along tectonically active as opposed to passive coasts (Johnson, 1988a), the Northpatagonian rocky shore represents an example of a flooded divergent margin.

### 5.3. Paleocology

Most of the data presented here reveals that the K/P transgressive bioclastic deposits are mainly composed of skeletal fragments of Oysters (Feruglio, 1936, 1949; Spalletti et al., 1993). In addition, oysters are the dominant molluscan group throughout almost all Patagonian basins for this time (del Río, 2021) and support a wide variety of boring and encrusting biota, including sponges, polychaetes, bivalves, fungi, algae, barnacles and bryozoans (Brezina et al., 2014, 2017). Though oysters colonized both hard and soft substrates (Seilacher et al., 1985; Machalski, 1998; Kidwell and Brenchley, 1994; Anderson et al., 2004), they are typically encrusting fauna of rocky shores from the Mesozoic (i.e., Johnson, 1988a; Zít and Nekvasilova, 1996; de Gibert et al., 1998,

2012; Johnson and Baarli, 1999, 2012; Betzler et al., 2000; D'Alessandro et al., 2004; Sørensen and Surlyk, 2008; Bover-Arnal et al., 2011; among others), even present in K/P rocky shore successions (i.e., Sanders, 1998). In this sense, Late Cretaceous carbonate platforms reached high latitudes in the northern and southern hemispheres (Kiessling et al., 2003), with benthic carbonate factories dominated by bivalves (oysters), echinoids, bryozoan, brachiopods, and red algae (i.e., Surlyk and Christensen, 1974; Surlyk, 1997). Oysters are considered as habitat-forming species or ecosystem engineers that enhance the biodiversity from attenuation wave energy, improvement of water quality by biofiltered, creation of habitats for fish species, and support of a wide diversity of epibenthic invertebrates (Newell, 1988; Coen et al., 2007; Grabowski and Peterson, 2007; Parras and Casadío, 2006; Padilla, 2010; Gutiérrez et al., 2011; Bayne, 2017; Coen and Humphries, 2017; McAfee et al., 2017). Therefore, the Northpatagonian rocky shore provided a regional ecological niche for oyster reefs and probably enhanced the K/P coastal biodiversity (Fig. 15).

The interplay of local/regional factors (substrate, paleorelief, subsidence) and external controls (sea level rise) favored the conformation of a K/P, short-lived, transgressive, cool-water carbonate factory (*sensu* Surlyk, 1997; Schlager, 2000, 2003; Reijmer, 2021) or heterozoan factory (*sensu* Michel et al., 2018, 2019) over the NPM. This example is partially synchronous (Late Cretaceous-Danian) with one of the largest and longest-lived cool-water carbonate platforms occurred in the Baltic Shield (Surlyk, 1997), mainly conformed by oyster bank communities along the rocky shorelines. The inclusion of the K/P cool water carbonate sedimentation over Patagonia in the worldwide carbonate platform reconstructions (i.e., Kiessling et al., 2000, 2003) will extend its distribution to higher latitudes and would increase the global carbonate production (70–140 km<sup>3</sup>).

## 6. Conclusions

- We recognized the K/P Northpatagonian rocky shore over 70.000 km<sup>2</sup> in central-northern extra-Andean Patagonia (Argentina), the widest known ancient rocky shore of South America. The substrate is mainly composed of Jurassic volcanic rocks (Chon Aike SLIP), with more restricted Triassic volcanic rocks and Paleozoic igneous-metamorphic complexes.

- The initial K/P rocky shore transgressive record is mainly composed of bioclastic deposits up to 40 m in thickness. The hard substrate, the irregular seacoast configuration, and the relatively low accommodation settings over the NPM provided the conditions for developing of oyster reefs and related marine biota.
- The interplay of geological, paleogeographical, and paleoecological conditions favored the conformation of a K/P, short-lived, transgressive, cool-water carbonate factory in the south-eastern extreme of South America.

## Data availability

Data will be made available on request.

## Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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